

*TC03 of ICG*  
*School Thermodynamic of Glass*

# Entropy, viscosity and structure of glasses and melts

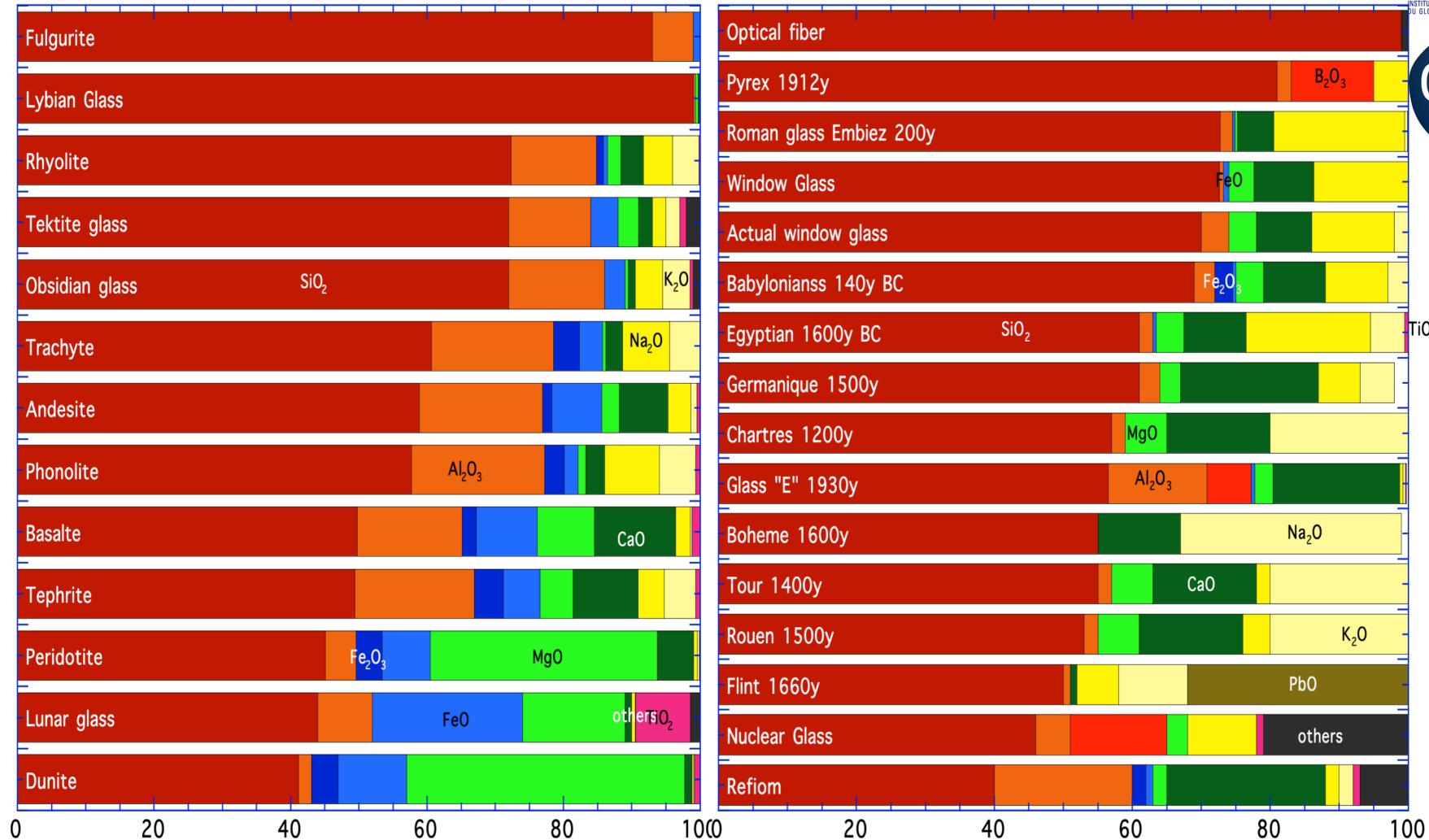
Daniel R. Neuville

Géomatériaux,  
CNRS-Institut de physique du globe de Paris,  
Université de Paris,  
1 rue Jussieu, F-75238 Paris cedex 05, France



# Huge similarities between natural glasses and human glasses....

Geomaterial



$$\mu_i = \Delta H_i - T \Delta S_{mi}$$

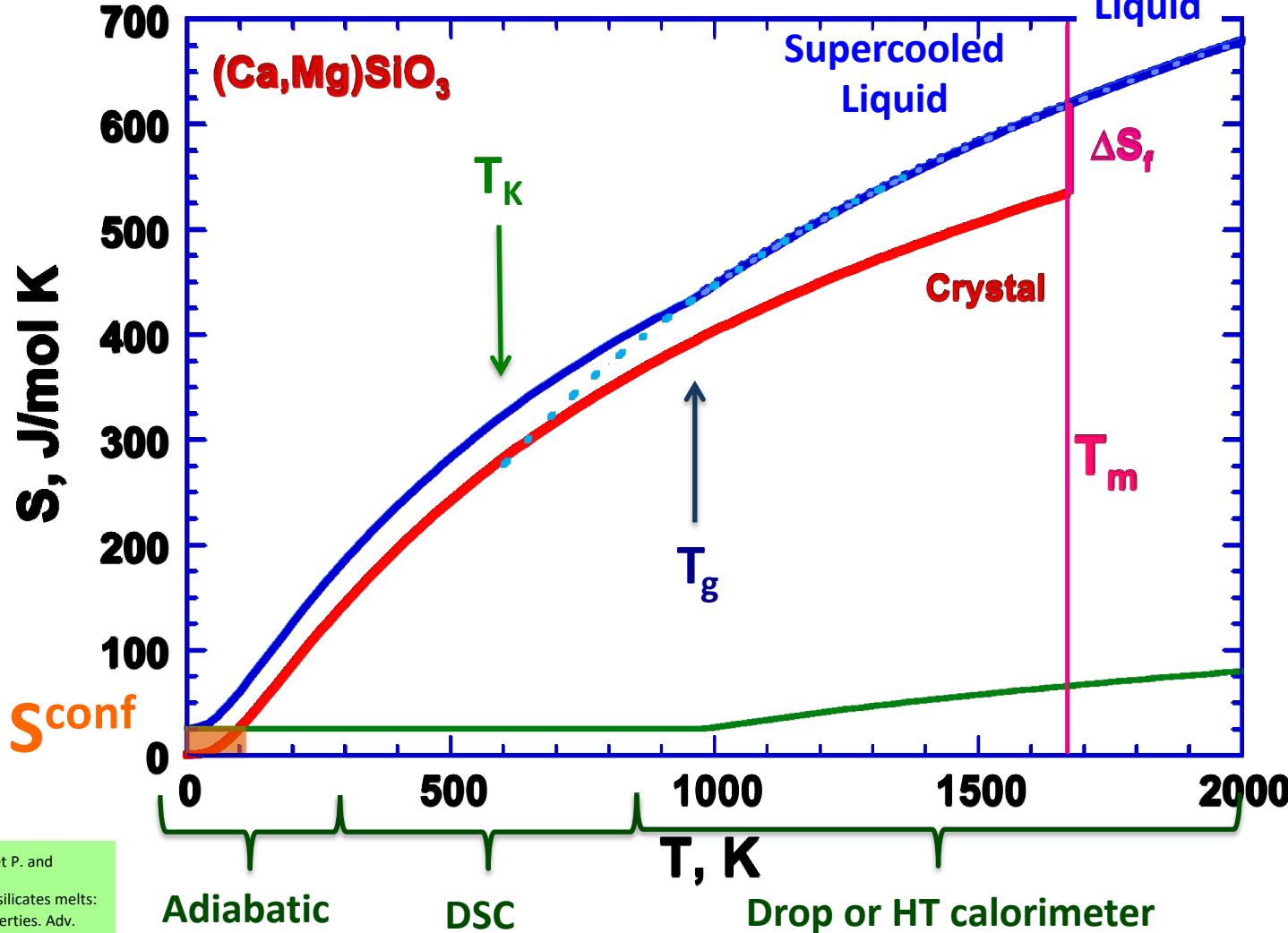
$$S^{\text{conf}}_{\text{melts}} \Rightarrow \Delta S_{\text{mix}} \Rightarrow \Delta S_{\text{mi}}$$



**Diopside**  
 **$\text{CaMgSi}_2\text{O}_6$**

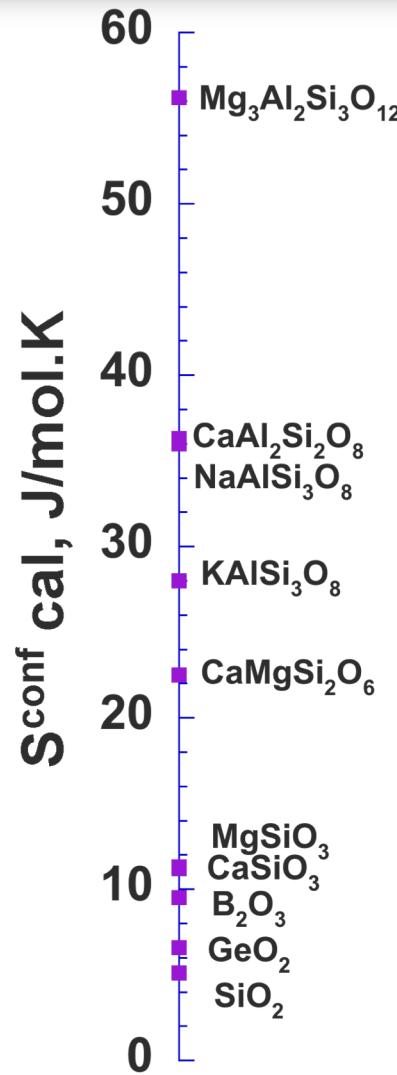


$$S^{conf}(Tg) = \int_0^{T_m} \frac{Cp^{Crystal}}{T} . dT + \Delta S_f + \int_{T_m}^{T_g} \frac{Cp^{liquid}}{T} . dT + \int_{T_g}^0 \frac{Cp^{glass}}{T} . dT$$

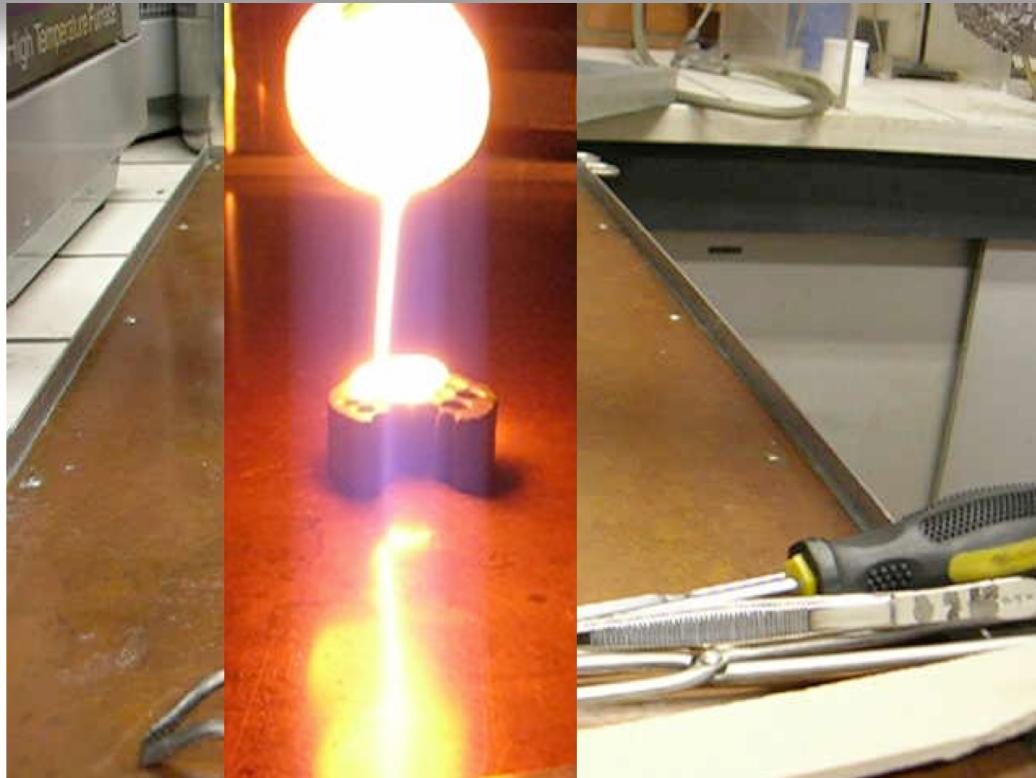


Modified from Richet P. and Neuville D.R. (1992)  
Thermodynamics of silicates melts:  
Configurational properties. Adv.  
Phys. Geochim., 10, 132-161.





Crystal



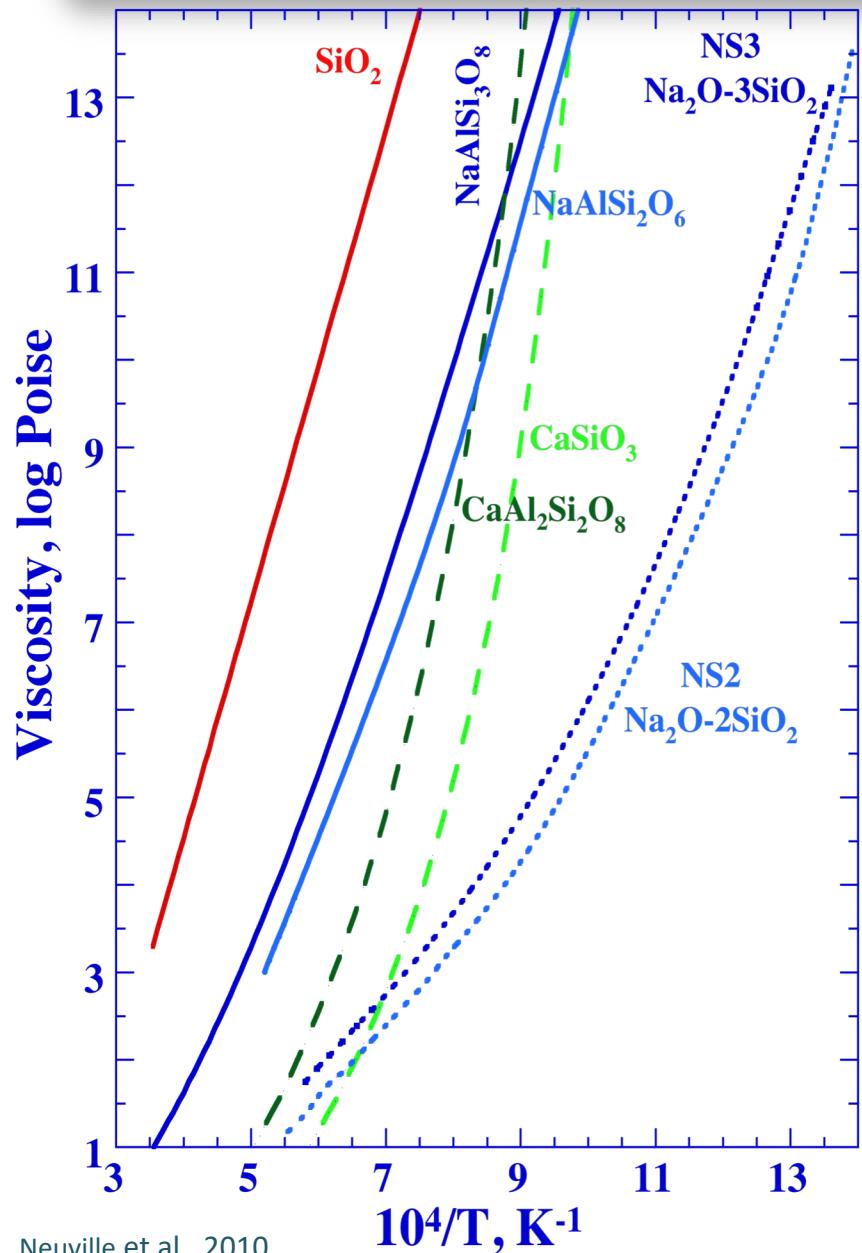
Glass



*Where is the crystal?*



Viscosity equation ?



Arrhenius :

$$\eta(T) = A \cdot \exp(E/RT)$$

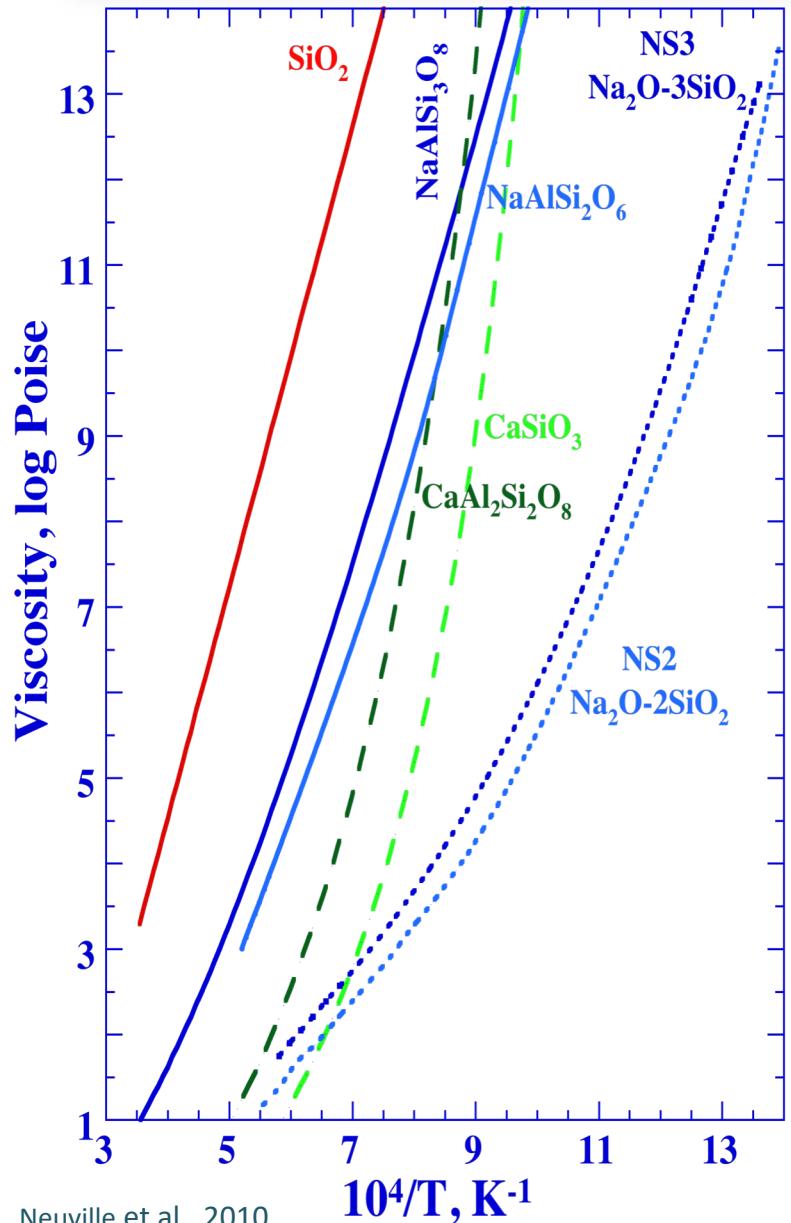
$$\Leftrightarrow \log \eta = A + B/T$$

Yes but only for  $\text{SiO}_2$ ,  $\text{GeO}_2$ ,  $\text{NaAlSiO}_8$ ,  $\text{KAISiO}_8$  because activation energy change from 2000 kJ/mol at 1000 K up down 300 kJ/mol at 1800 K for NS3.

Need TVF equation  
 $\log \eta = A_1 + B_1/(T-T_1)$

But, just a fit .....





$$\eta(T) = A_e \cdot \exp[B_e / TS^{\text{conf}}(T)]$$

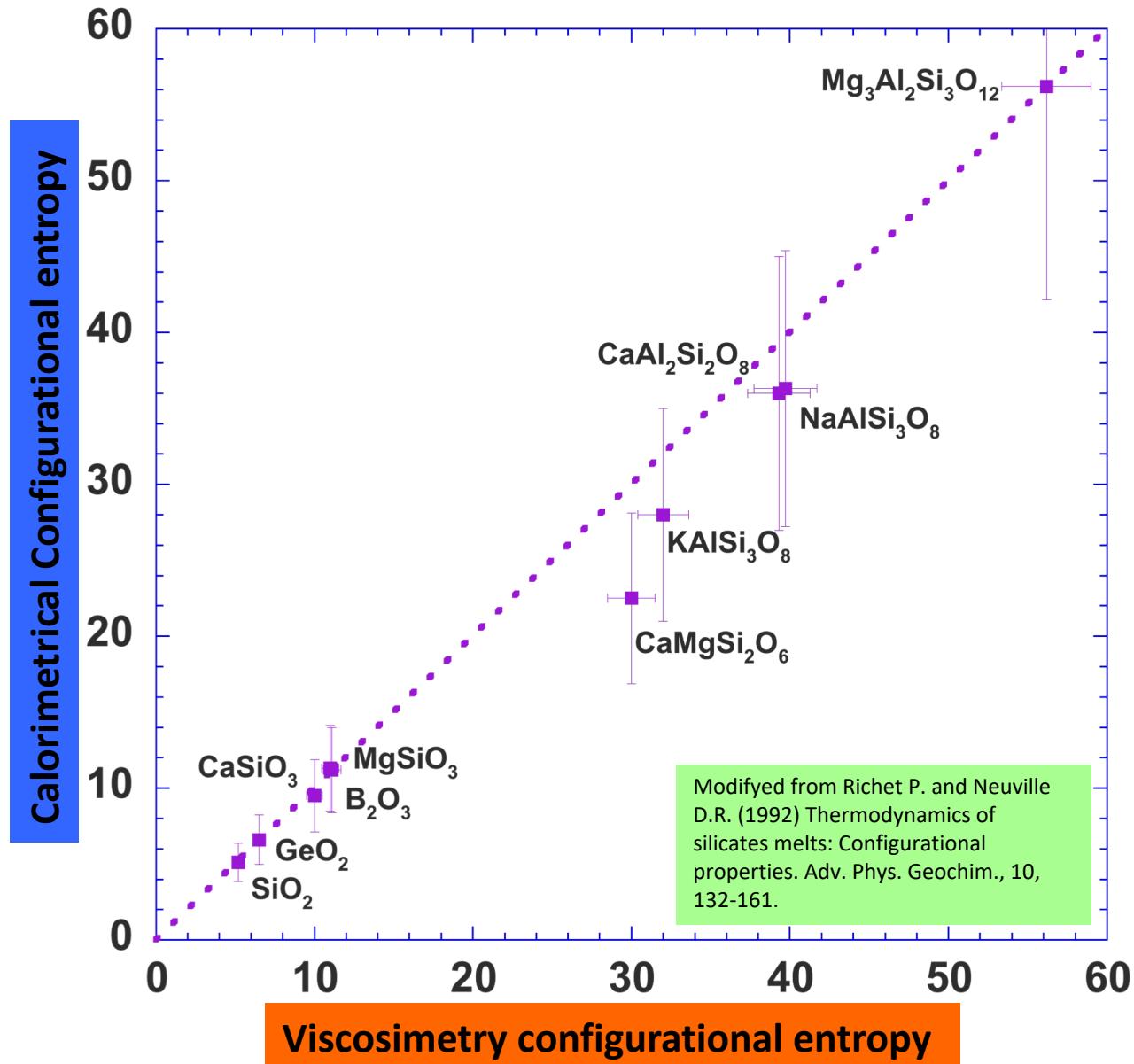
**Proposed by Adam and Gibbs, 1964**  
**First used to silicate melts by Urbain, 1972,**  
**Wong and Angell 1976,**  
**Scherer, 1984, Richet, 1984, ...**  
**Neuville and Richet, 1991....**

$$S^{\text{conf}}(T) = S^{\text{conf}}(T_g) + \int_{T_g}^T Cp^{\text{conf}} / T dt$$

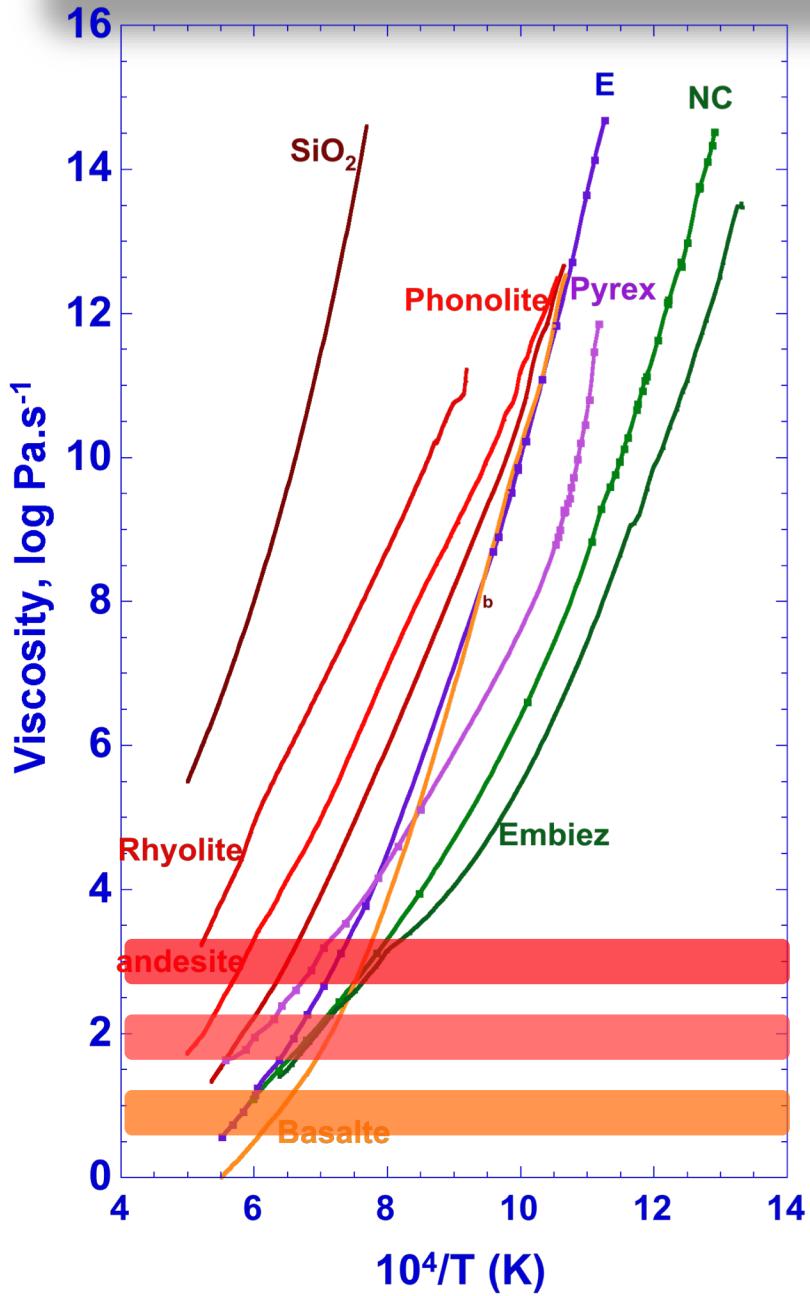
$$Cp^{\text{conf}}(T) = Cpg(Tg) - Cpl(T)$$

**Calorimetry measurements**  
**=> Easy**





## What is viscosity ?



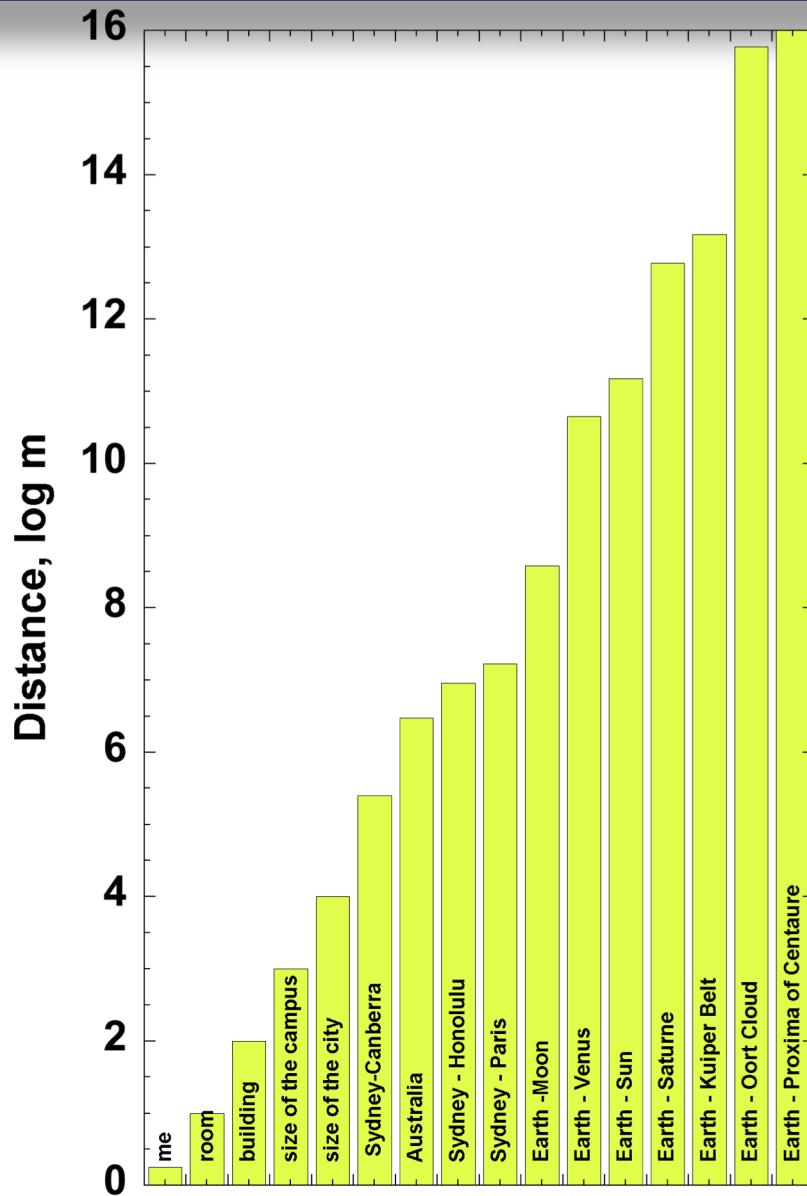
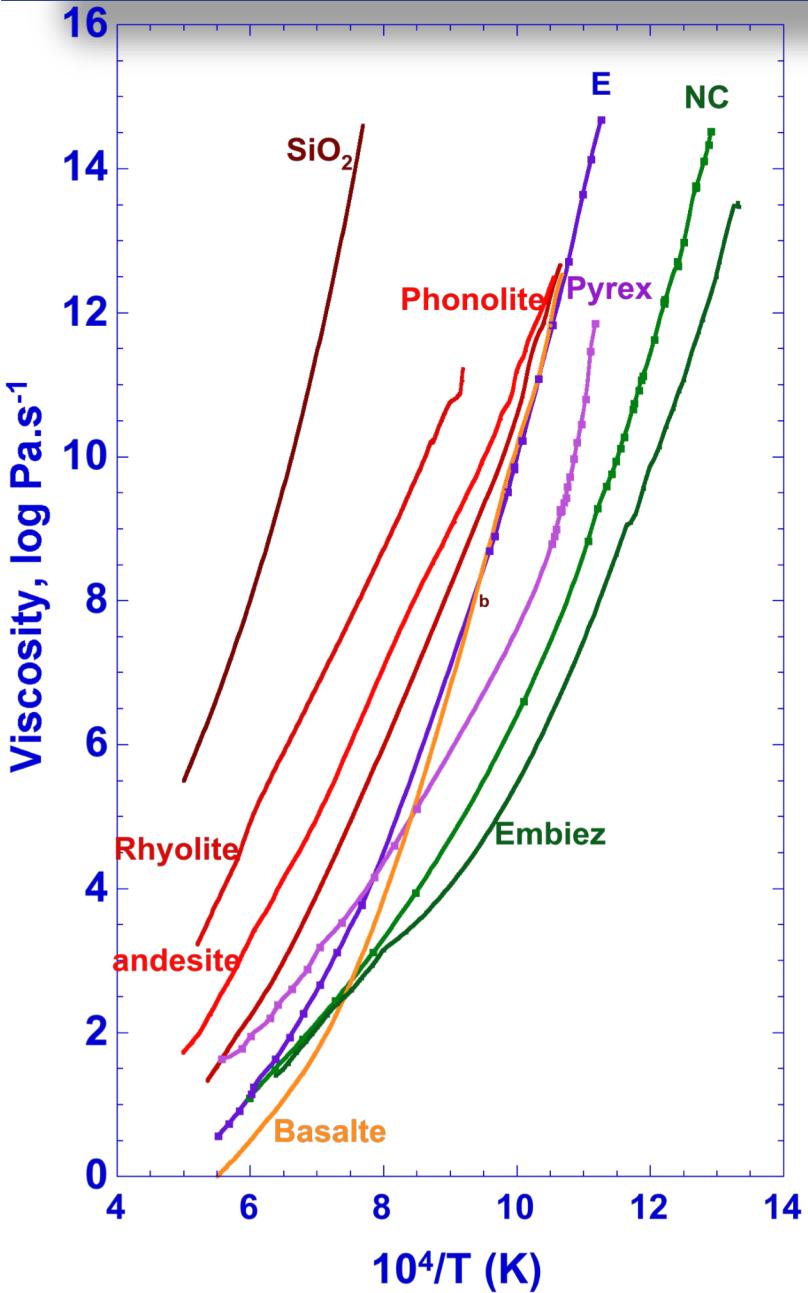
1

2

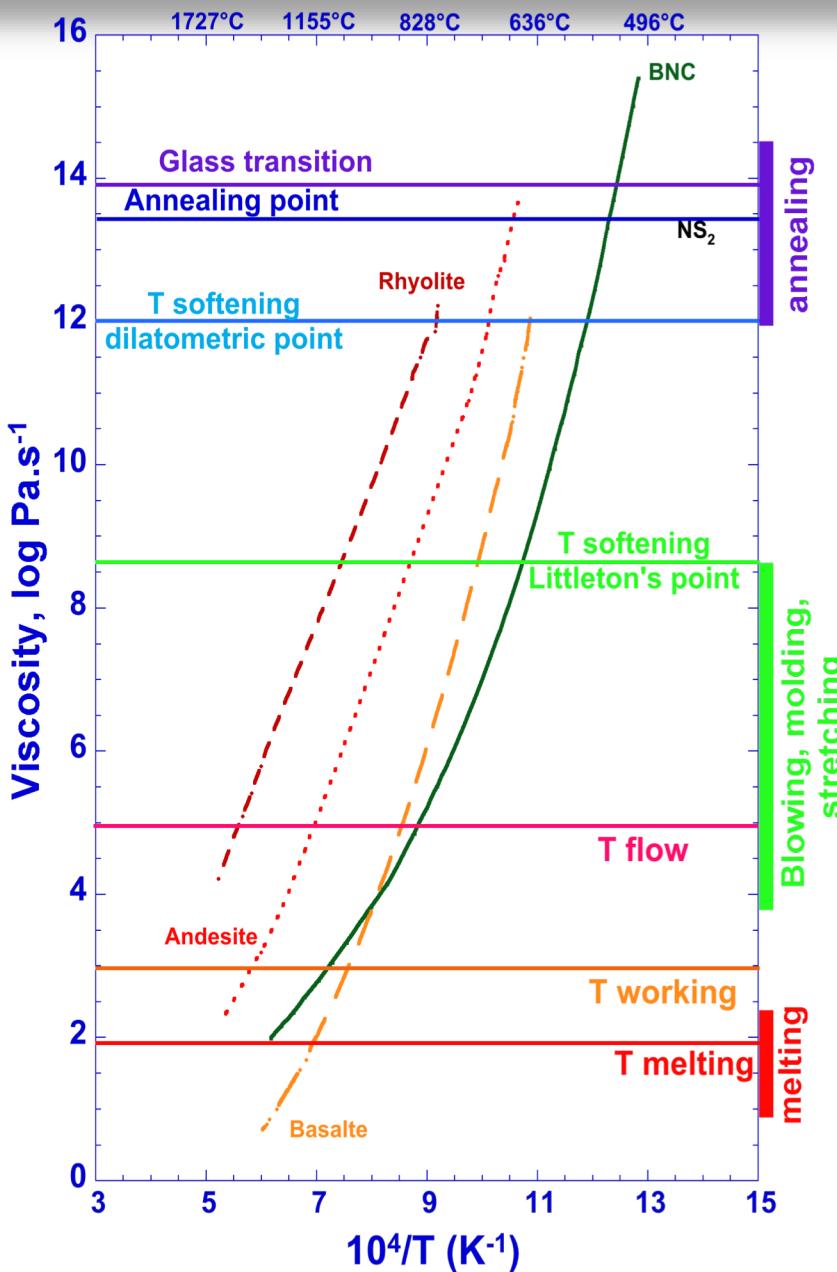
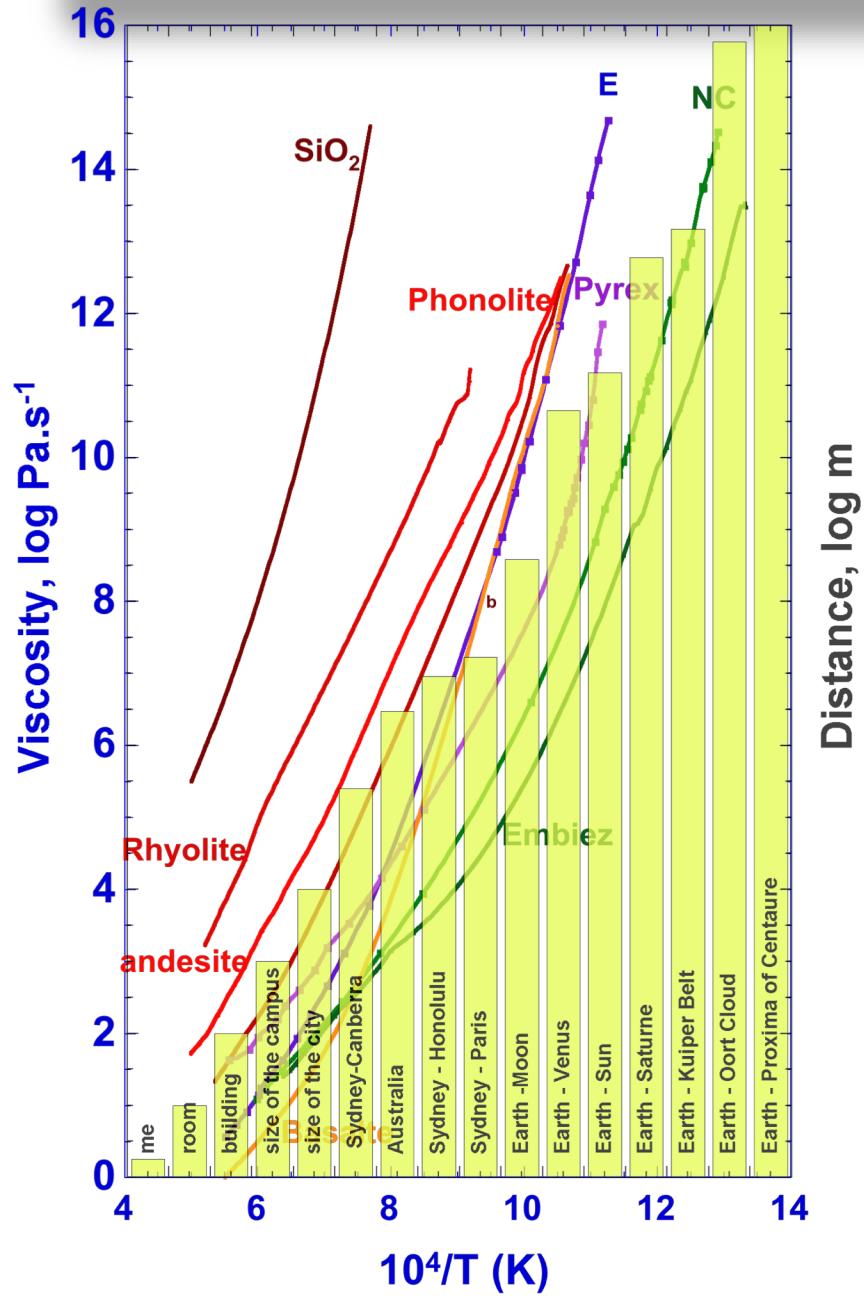
3

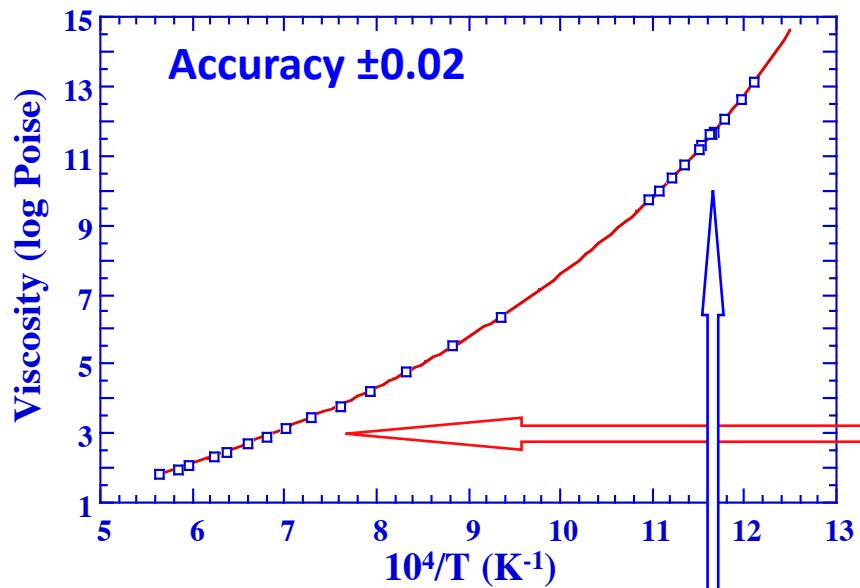


## What is viscosity ?

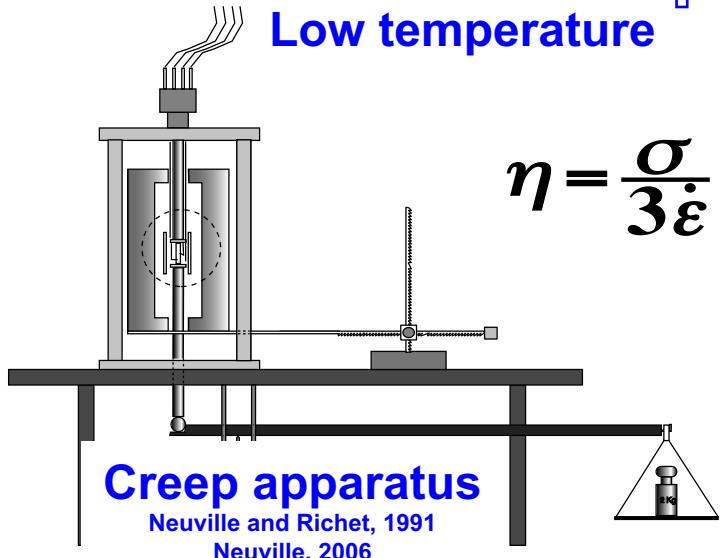


## What is viscosity ?

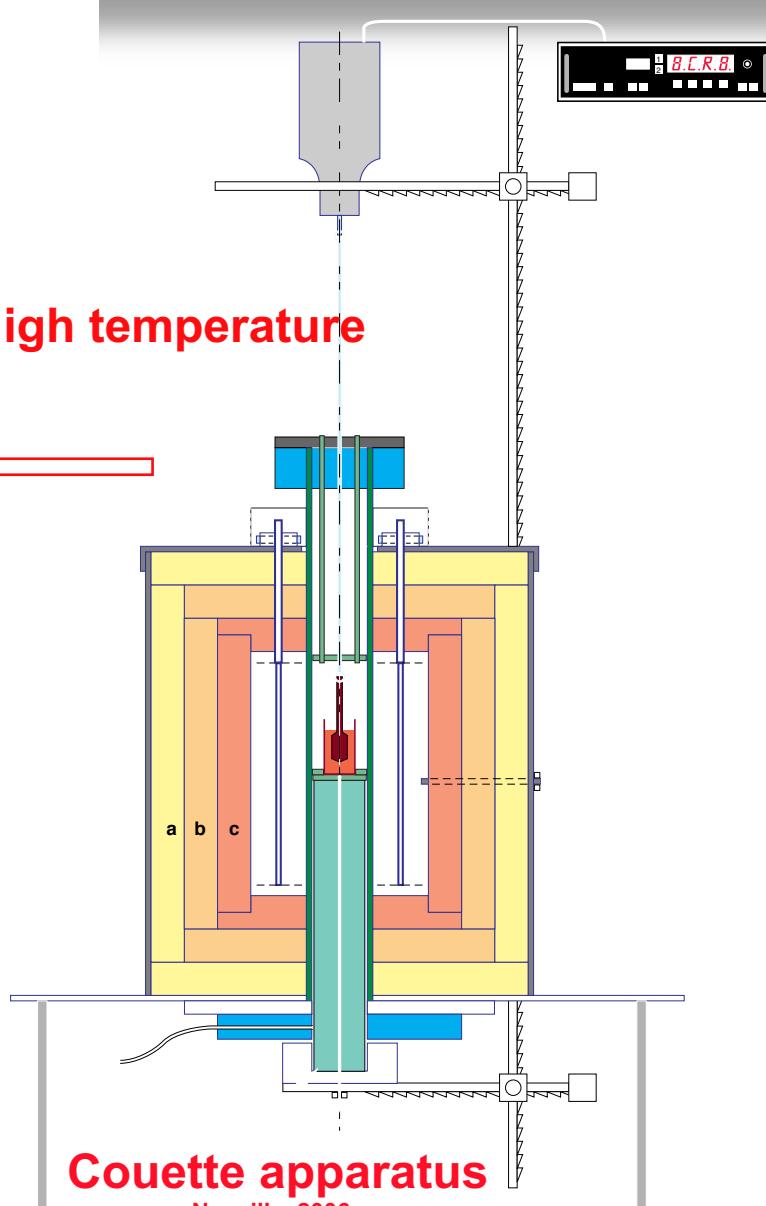




High temperature

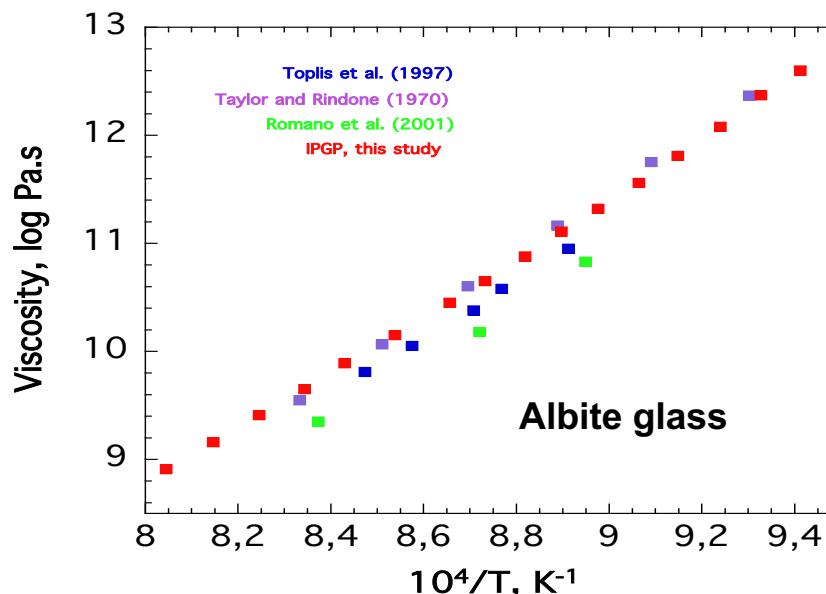
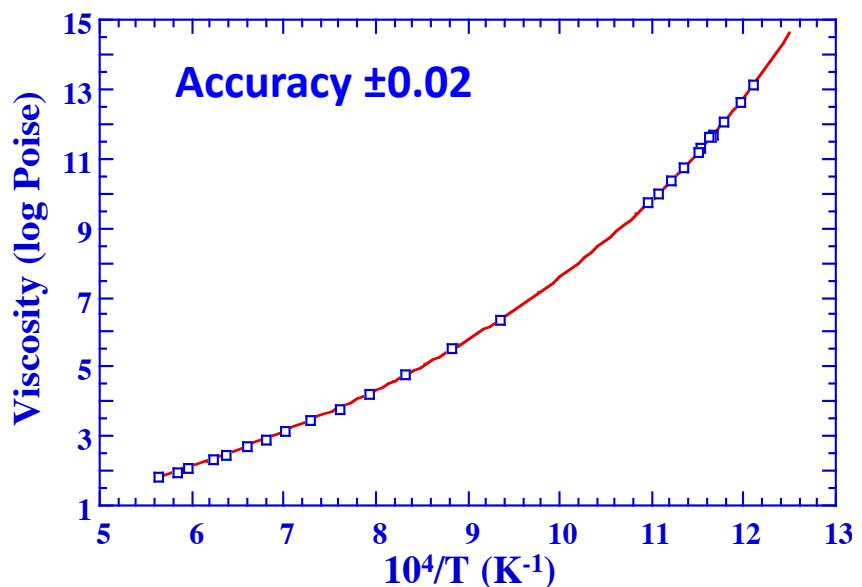
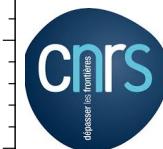


$$\eta = \frac{\sigma}{3\dot{\epsilon}}$$

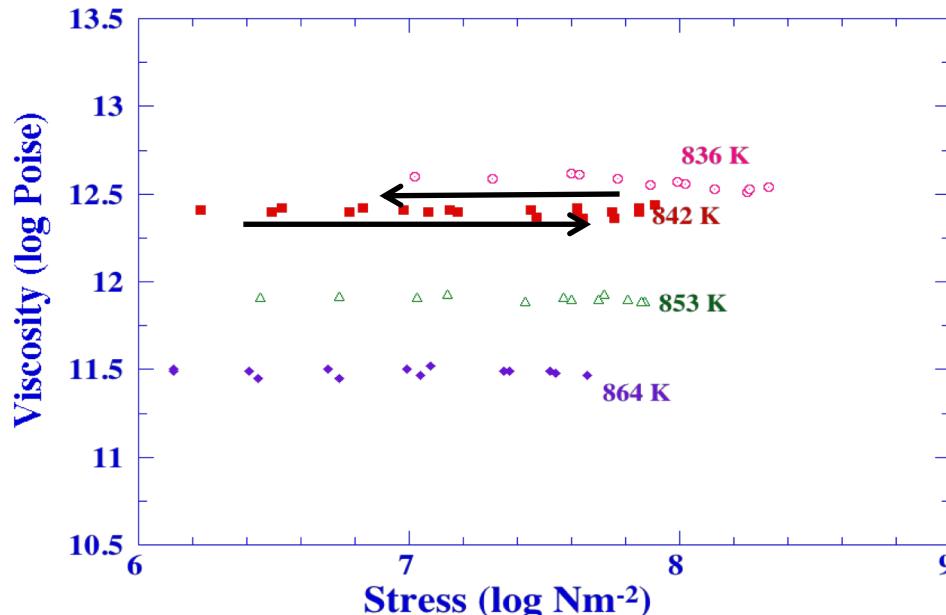


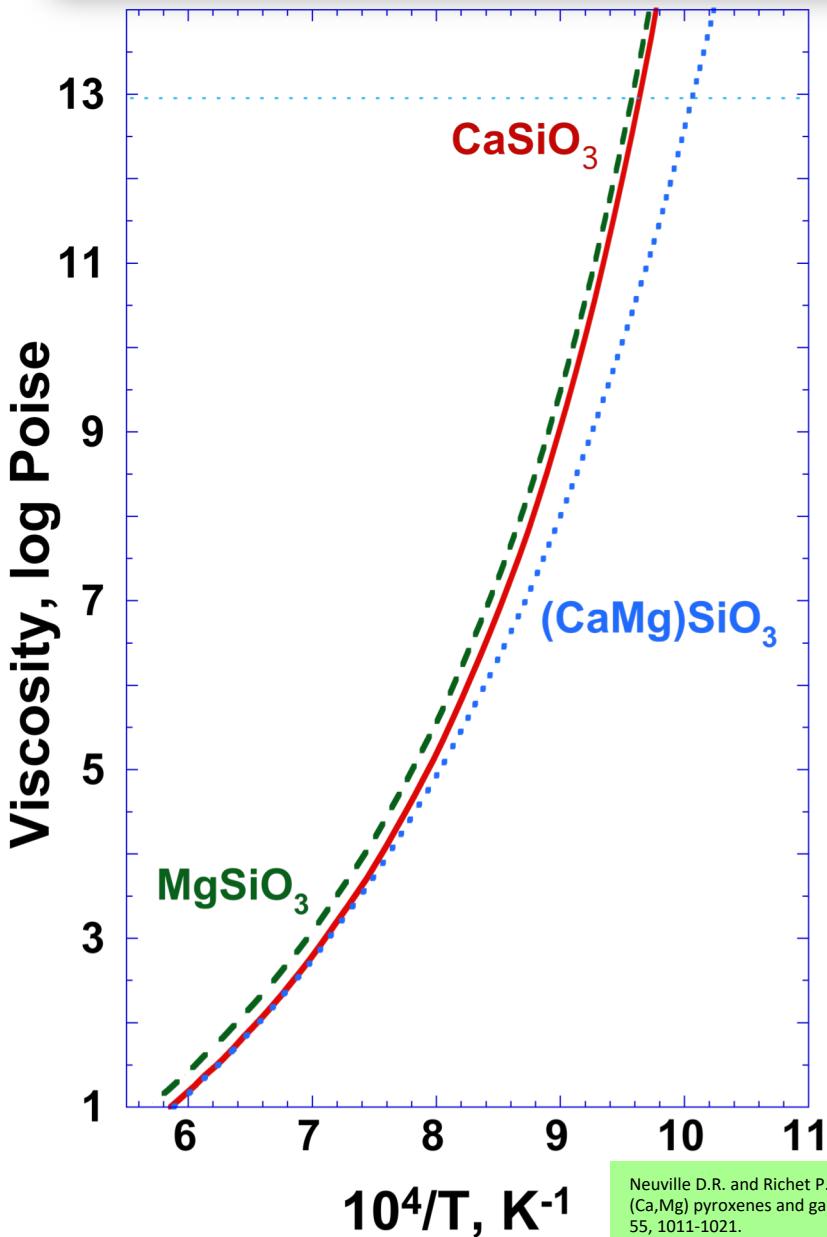
# Viscosity measurements

## Geomaterial

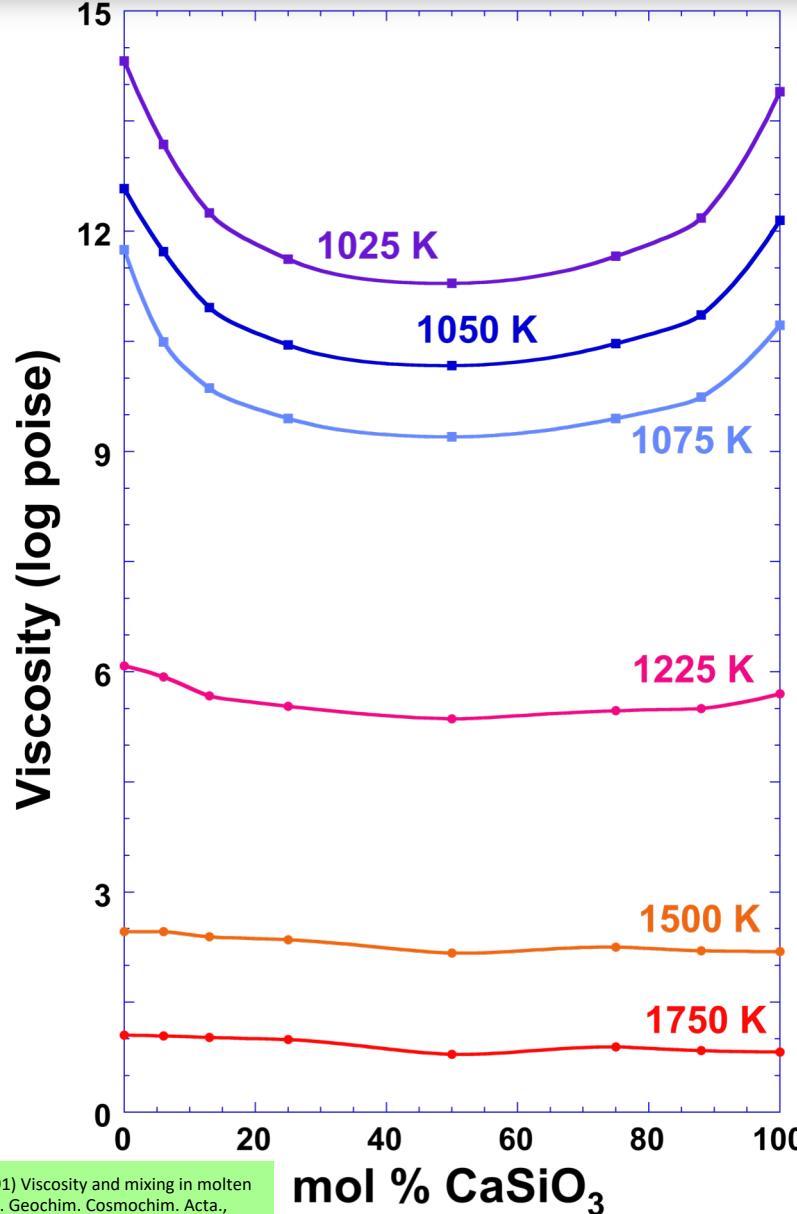


**Liquid =>  
Newtonian behavior**





Neuville D.R. and Richet P. (1991) Viscosity and mixing in molten (Ca,Mg) pyroxenes and garnets. Geochim. Cosmochim. Acta., 55, 1011-1021.



$$\log \eta = A_e + B_e / TS^{\text{conf}}(T)$$

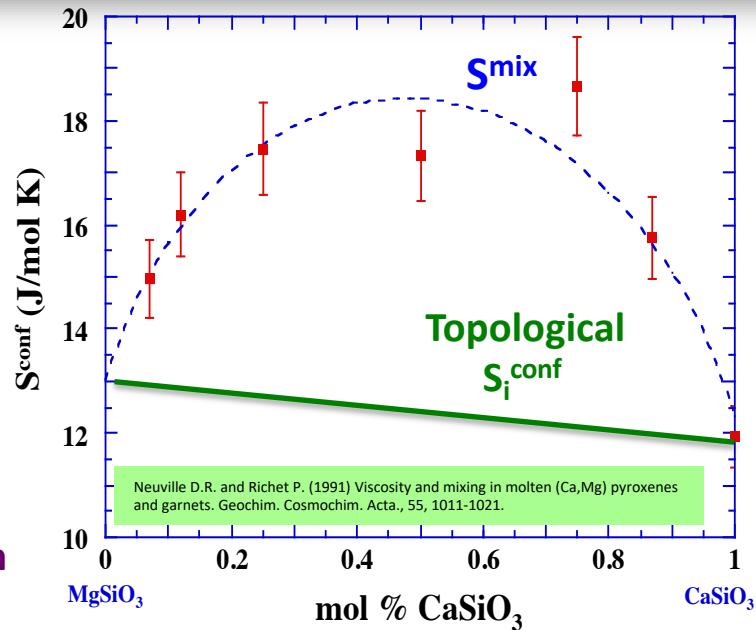
$$S^{\text{conf}}(T) = S^{\text{conf}}(T_g) + \int_{T_g}^T Cp^{\text{conf}} / T dt$$

$$C_p^{\text{conf}} = C_p^{\text{l}} - C_{\text{pg}}(T_g)$$

$$S^{\text{conf}}(T_g) = S^{\text{mix}} + \sum X_i S_i^{\text{conf}}(T_g)$$

$$S^{\text{mix}} = -nR \sum X_i \ln X_i \quad \text{with } X_i = \text{Ca}/(\text{Ca}+\text{Mg})$$

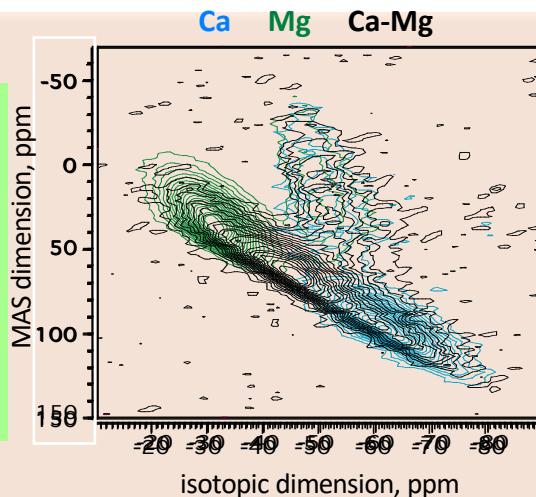
Ideal mixing => random distribution



### O-NMR

Allwardt and Stebbins 2004

- “viewpoint” of the NBO
- <sup>17</sup>O chemical shifts depend strongly on which cations are nearby



- detailed analyses of spectra support almost random distribution of Ca + Mg around NBO
- size difference of Ca<sup>2+</sup> and Mg<sup>2+</sup> is insufficient to cause ordering

Configurational entropy : image of the structure of the liquid

# Ca/Na mixing

# Geomaterial

Morey and Bowen, (1925)

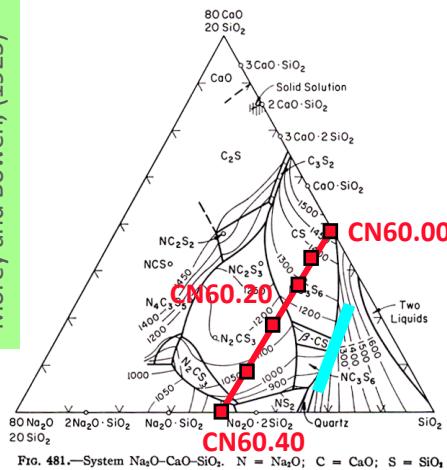
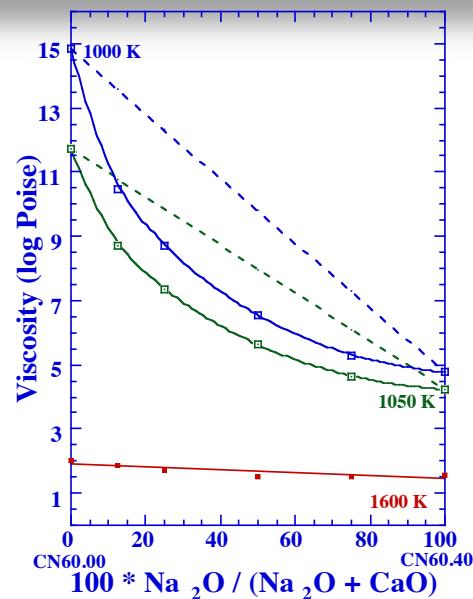
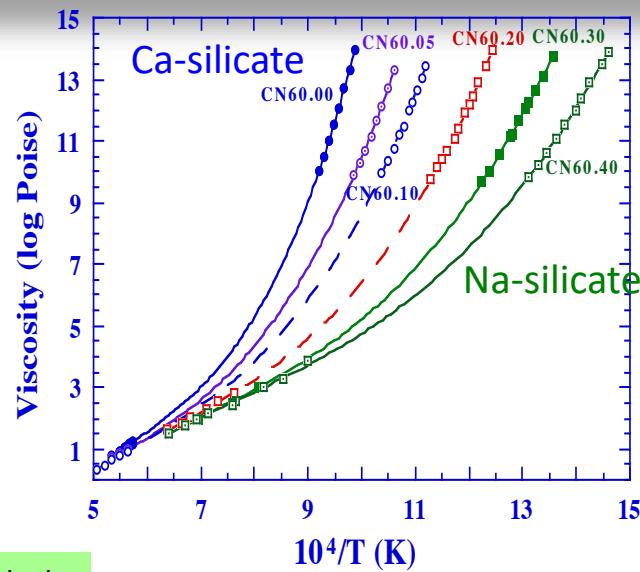
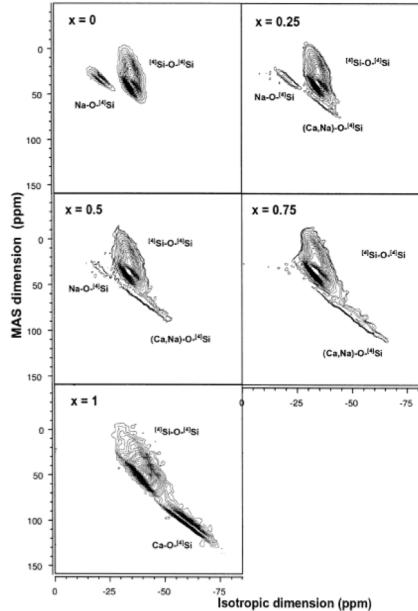
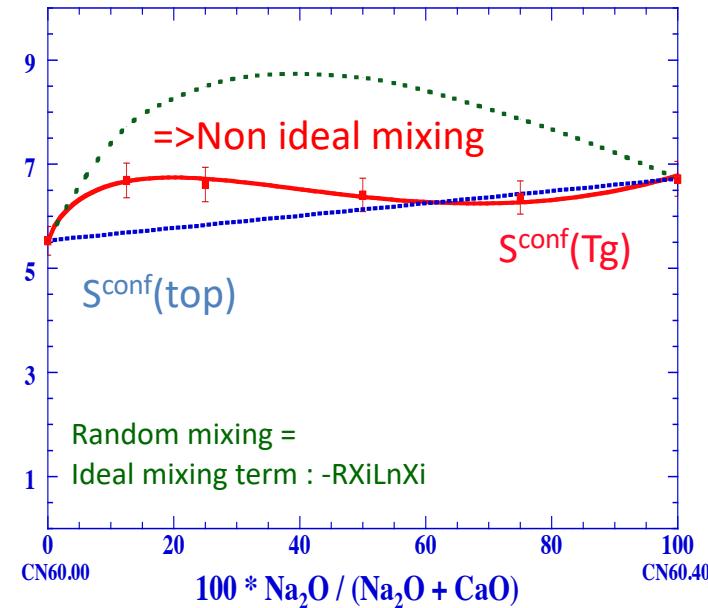


FIG. 481.—System  $\text{Na}_2\text{O}-\text{CaO}-\text{SiO}_2$ . N =  $\text{Na}_2\text{O}$ ; C =  $\text{CaO}$ ; S =  $\text{SiO}_2$ . The area NS-CS-SiO<sub>2</sub> after Morey and Bowen, Fig. 482.  
E. R. Segnit, Am. J. Sci., 251 [8] 590 (1958).



Neuville D.R. (2006) Viscosity, structure and mixing in (Ca, Na) silicate melts. Chem. Geol., 229, 28-42.



Raman spectroscopy (Neuville, 2006) and <sup>17</sup>O NMR (Lee and Stebbins, 2003) show a **non random distribution of Na and Ca**.

$A_e$  pre-exponential term

Viscosity at “infinite” temperature

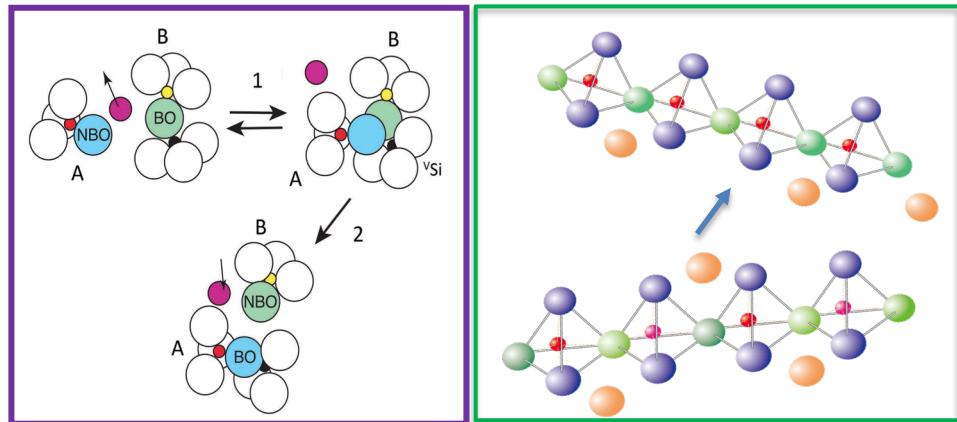
$$\log(\eta) = A_e + \frac{B_e}{T S^{conf}(T)}$$

Viscous flow occurs through cooperative rearrangements of subunits

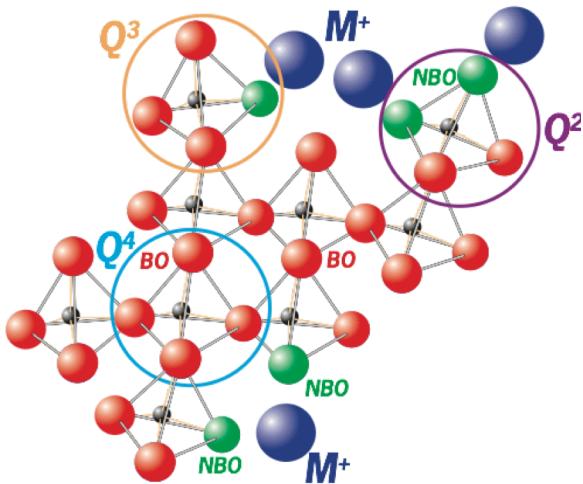
$B_e$  proportional to the energy barriers  $\Delta\mu$  to viscous flow

$$S^{conf}(T_g) + \int_{T_g}^T \frac{C_p^{conf}}{T} dt$$

*In silicate melts, viscous flow is related to exchange of oxygen by and interactions between high coordinated units (Farnan and Stebbins, 1992) or depolymerized species (Neuville et al., 2008)*



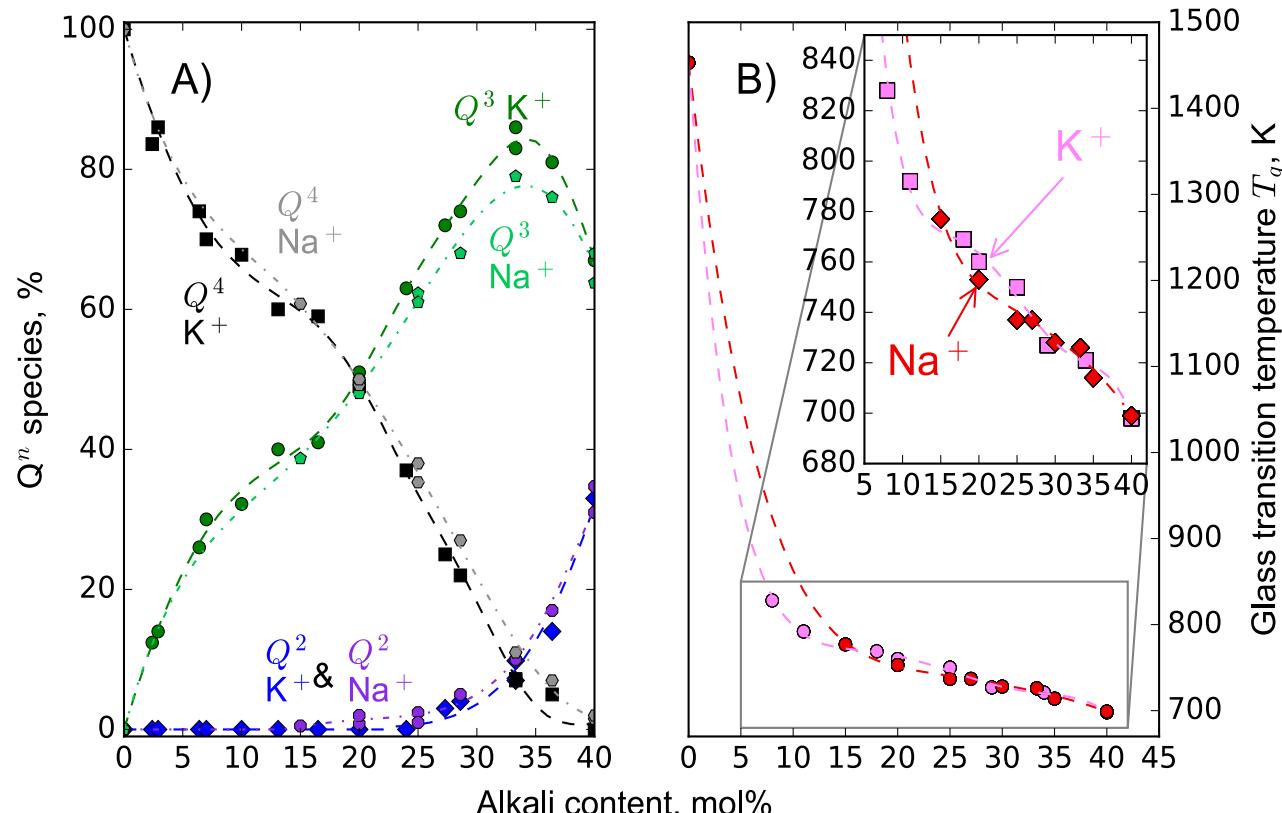
- glass and liquid heat capacity (e.g., Richet and Bottinga 1985, Richet 1987)
- glass molecular structure (e.g. from  $^{29}\text{Si}$  NMR: Maekawa et al. 1991, Sen and Yougman 2003)
- liquid viscosity (e.g., Bockris 1954, Poole 1949, Neuville 1992, Neuville 2006)



BO: bridging oxygen

NBO: non-bridging oxygen

$\text{M}^+$ : alkali cation



$$\log(\eta) = \cancel{A_e} + \frac{\cancel{B_e}}{T \left( S^{conf}(T_g) + \int_{T_g}^T \cancel{\frac{C_p^{conf}}{T}} dT \right)}$$

$B_e$  and  $S^{conf}(T_g)$  = sum of partial  $B_e$  and  $S^{conf}(T_g)$  molar values for each tetrahedral species

$\text{Na}_2\text{SiO}_3$  and  $\text{K}_2\text{SiO}_3$  ( $\text{Na}-\text{Q}^2$  and  $\text{K}-\text{Q}^2$ )

$\text{Na}_2\text{Si}_2\text{O}_5$  and  $\text{K}_2\text{Si}_2\text{O}_5$  ( $\text{Na}-\text{Q}^3$  and  $\text{K}-\text{Q}^3$ )

$\text{SiO}_2$  ( $\text{Q}^4$ )

Need to add to  $S^{conf}(T_g)$ :

- ideal mixing of Si between  $\text{Q}^2$ ,  $\text{Q}^3$  and  $\text{Q}^4$  units
- ideal mixing of Na and K in network modifier channels

Need to add to  $B_e$ :

- the above ideal mixing terms, scaled by two coefficients

- 13 parameters, 326 data points, from 60 to 100 mol% SiO<sub>2</sub>
- Minimisation of the least-square criterion using the Ipopt (Wächter and Biegler 2006) and Knitro (Byrd et al. 2006) solvers, and error calculation via non-parametric bootstrapping (20 000 re-sampling done)

*Partial S<sup>conf</sup>(T<sub>g</sub>),  
J mol<sup>-1</sup> K<sup>-1</sup>*

Na-Q<sup>2</sup>: 5.8 [4.8 - 6.9]

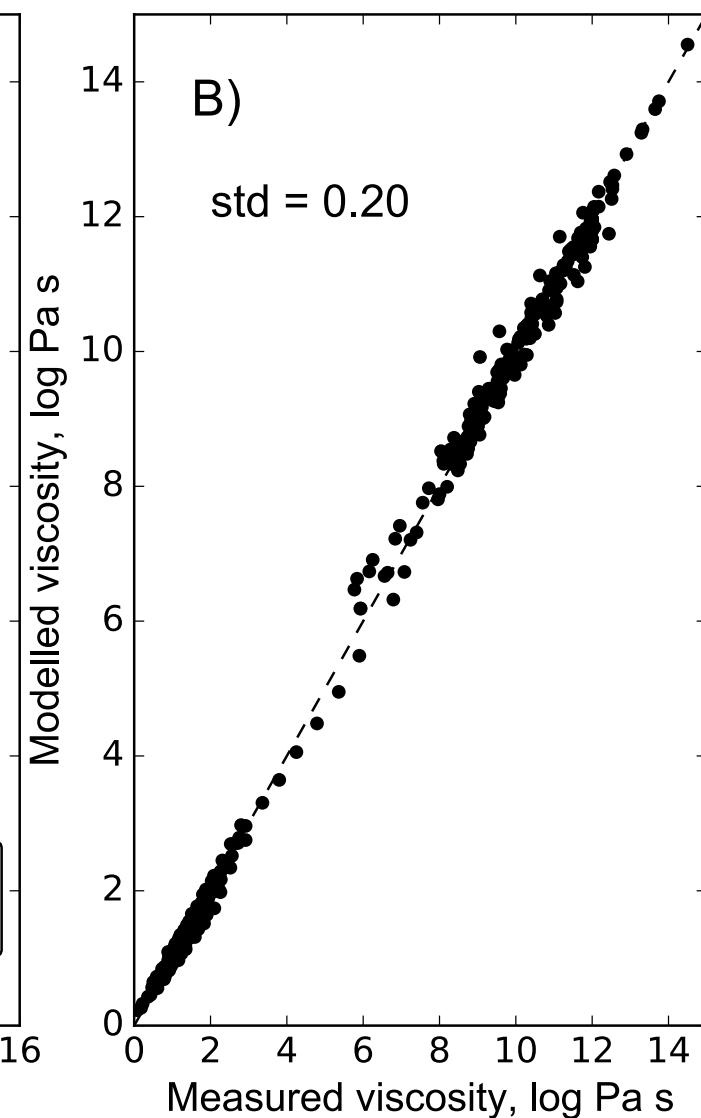
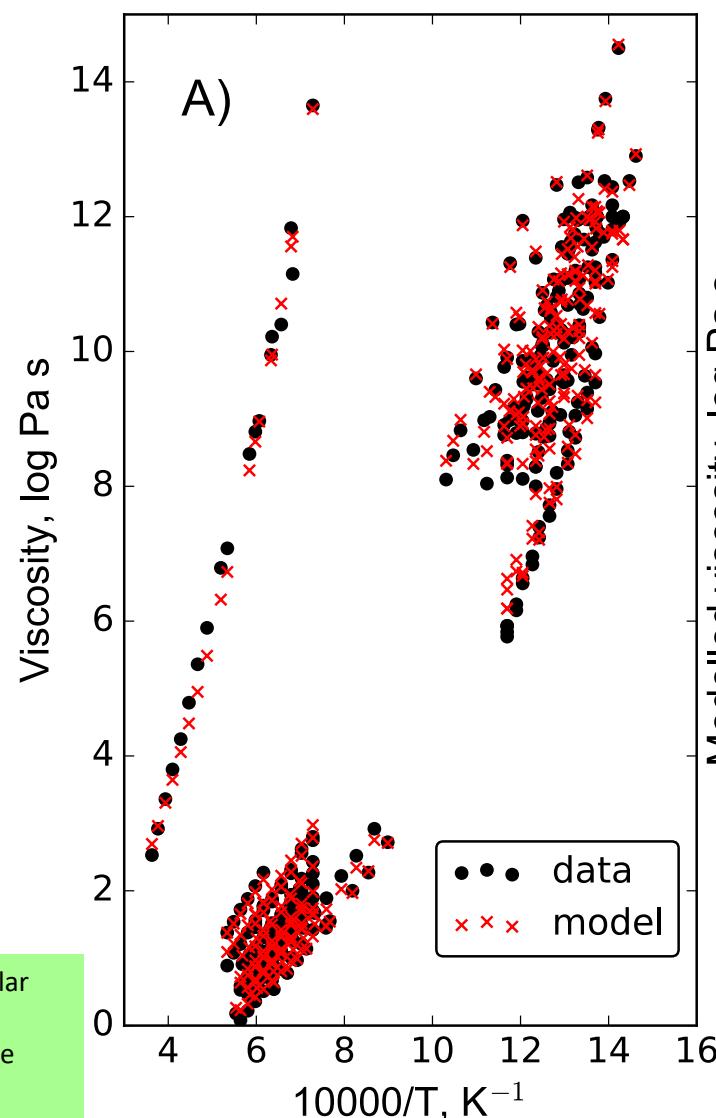
Na-Q<sup>3</sup>: 0.9 [0.3 - 1.3]

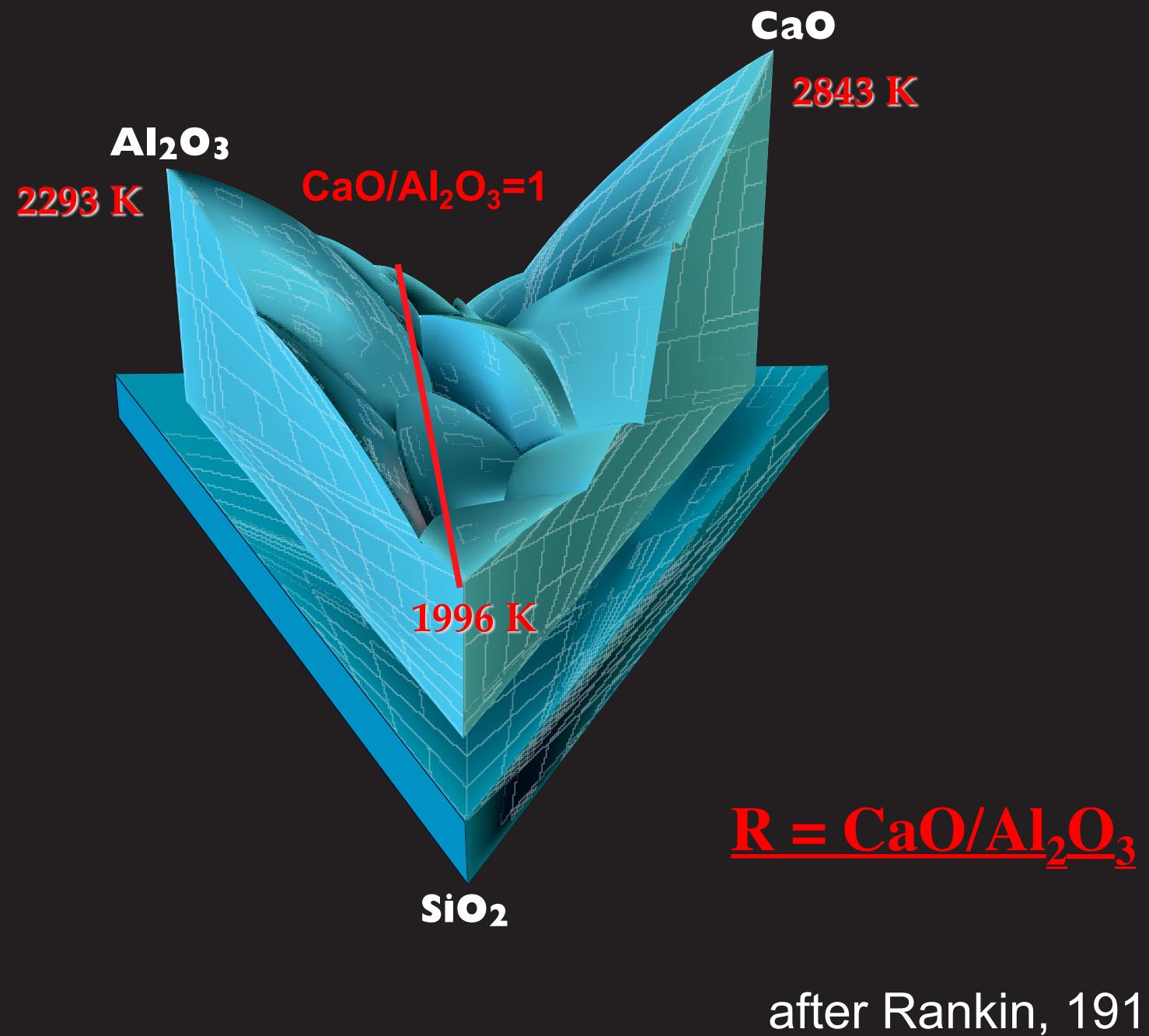
K-Q<sup>2</sup>: 12.0 [6.8 - 20.4]

K-Q<sup>3</sup>: 4.4 [3.3 - 5.4]

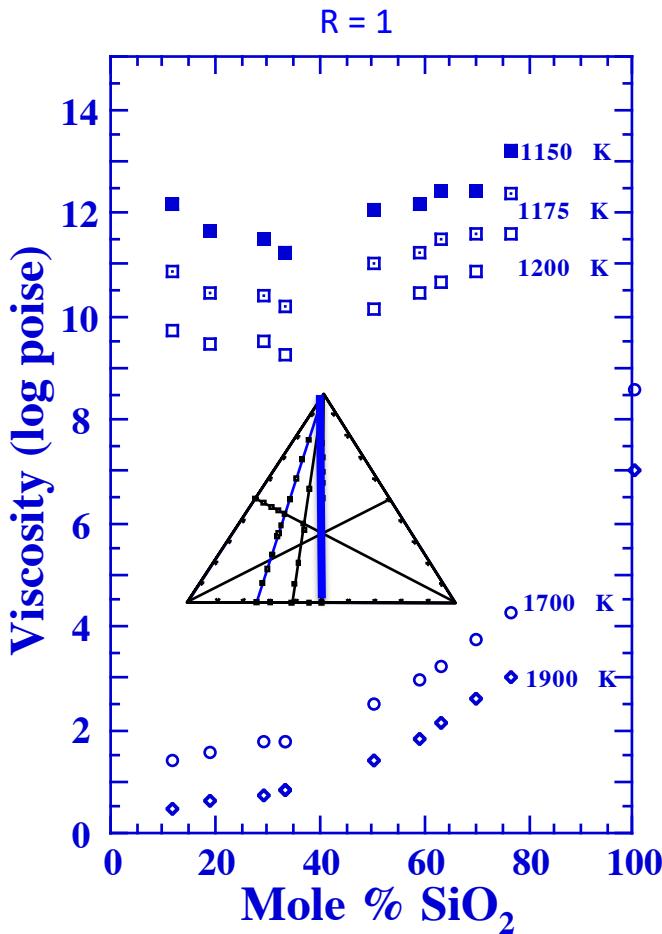
Q<sup>4</sup>: 7.9 [7.1 - 8.4]

SiO<sub>2</sub> exp. is 5.1 ± 2





# Configuration Entropy Theory



$$R = \text{CaO}/\text{Al}_2\text{O}_3$$

$$\log \eta = A_e + B_e/TS^{\text{conf}}(T)$$

$$S^{\text{conf}}(T_g) = S^{\text{mix}} + S^{\text{conf top}}$$

$$S^{\text{conf top}} = \sum x_i S_i^{\text{conf}}(T_g).$$

$$S^{\text{mix}} = -nR \sum X_i \ln X_i$$

$$X_i = \text{Al}/(\text{Al}+\text{Si})$$

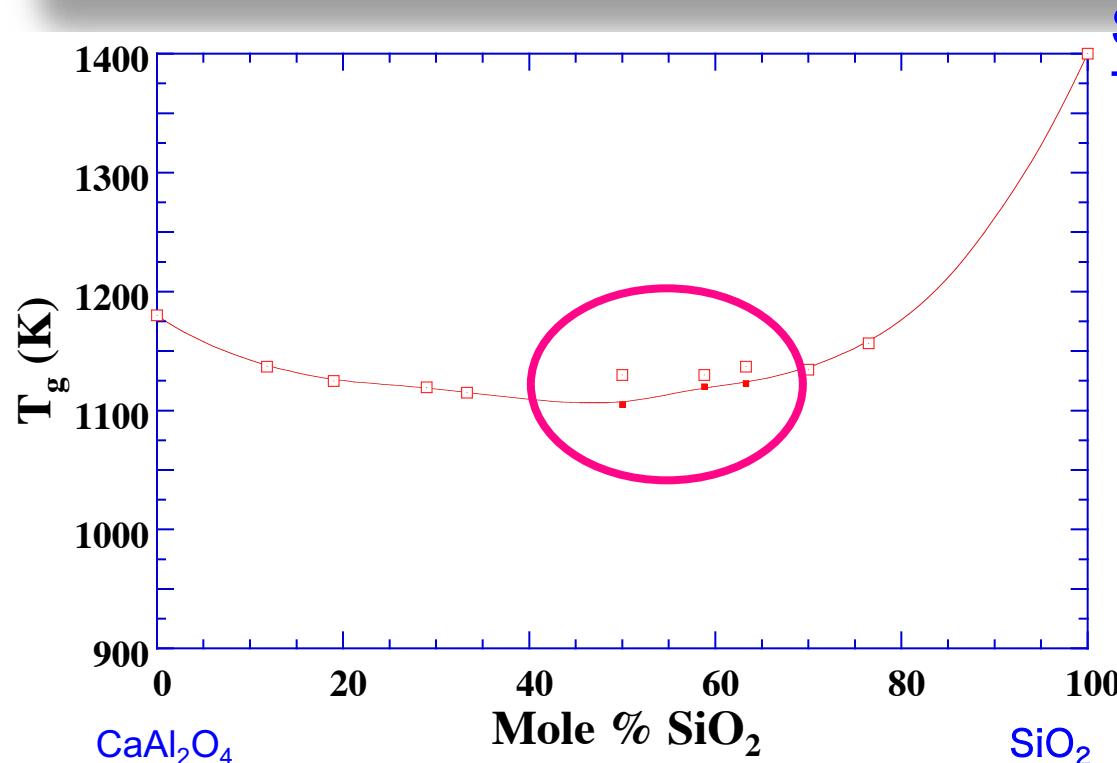
Ideal mixing  $\Rightarrow$  random distribution

$$S^{\text{conf}}(T) = S^{\text{conf}}(T_g) + \int_{T_g}^T C_p^{\text{conf}}/T dT$$

$$\text{with } C_p^{\text{conf}} = C_p^I - C_{pg}(T_g)$$

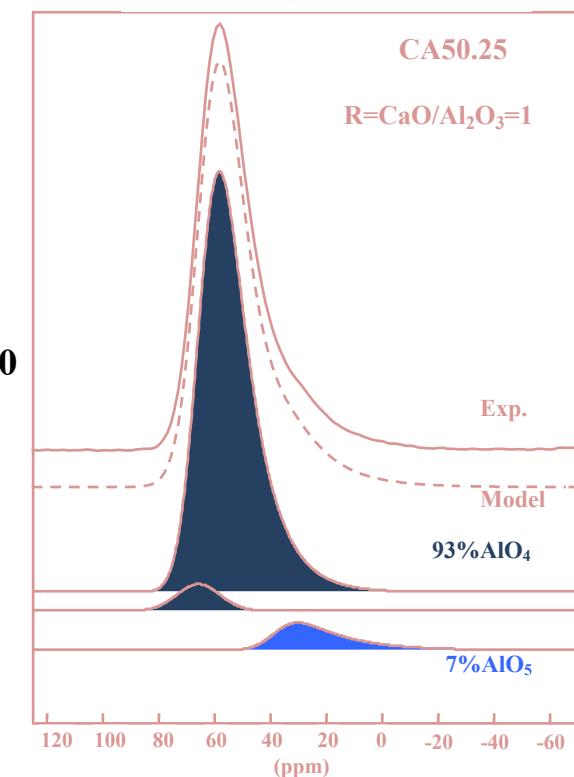
Substitution of Si by Al  
in Q<sup>4</sup> species along the join R=1





$\text{SiO}_2$   
 $T_g=1400\text{K}$

RMN 750MHz, CRMHT,  
Orléans,  $^{27}\text{Al}$  1D MAS



$\text{SiO}_2 \Rightarrow$  Tetrahedra  $\text{SiO}_4$

$\text{CaAl}_2\text{O}_4 \Rightarrow$  Tetrahedra  $\text{AlO}_4$

substitution of 1 Si by 1 Al and  
Ca charge compensator

Neuville D.R., Cormier L. and Massiot D. (2004) Role of aluminium in peraluminous region in the CAS system. Geochim. Cosmochim. Acta., 68, 5071-5079



# $\text{Na}_2\text{O}-\text{Al}_2\text{O}_3-\text{SiO}_2$ system

NA75.X : 75%  $\text{SiO}_2$

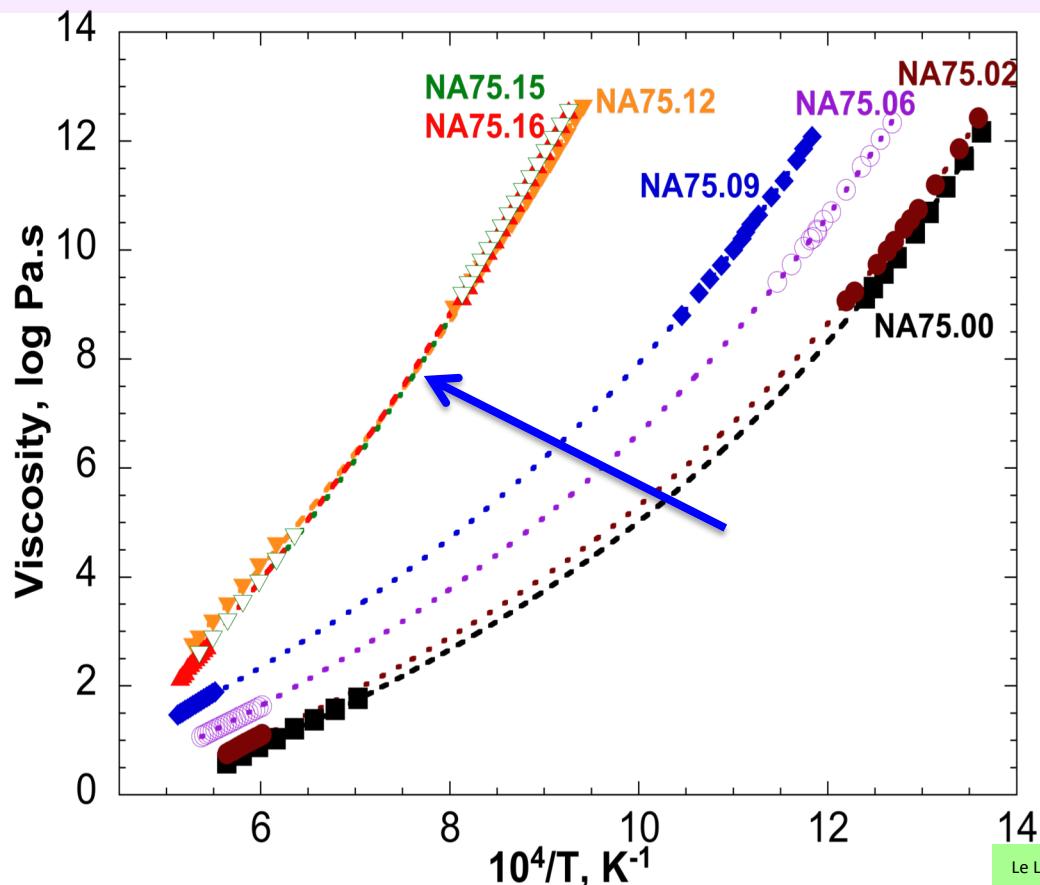
X=% $\text{Al}_2\text{O}_3$

$\text{Na}_2\text{O}=100-(75+\text{X})$

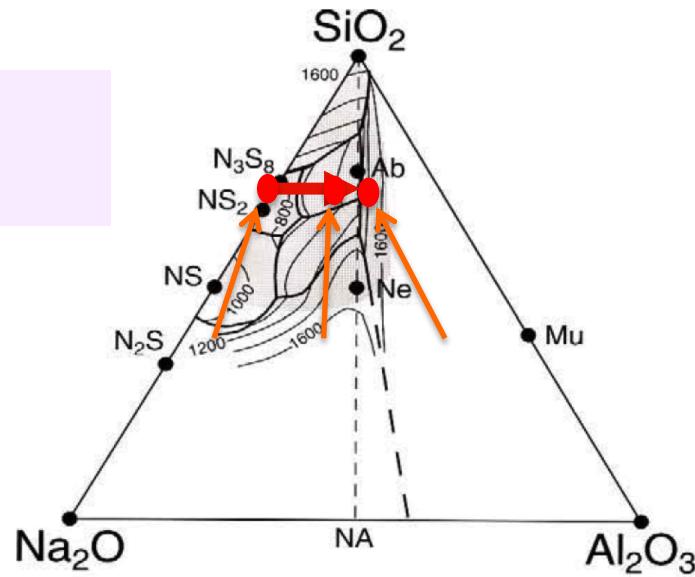
High-viscosity measurements  $10^8\text{-}10^{14}\text{ Pa.s} \pm 0.02$ : Creep apparatus

(Neuville and Richet, 1991, Neuville 2006)

Low-viscosity measurements  $1\text{-}10^5\text{ Pa.s} \pm 0.04$ : (Neuville, 2006)



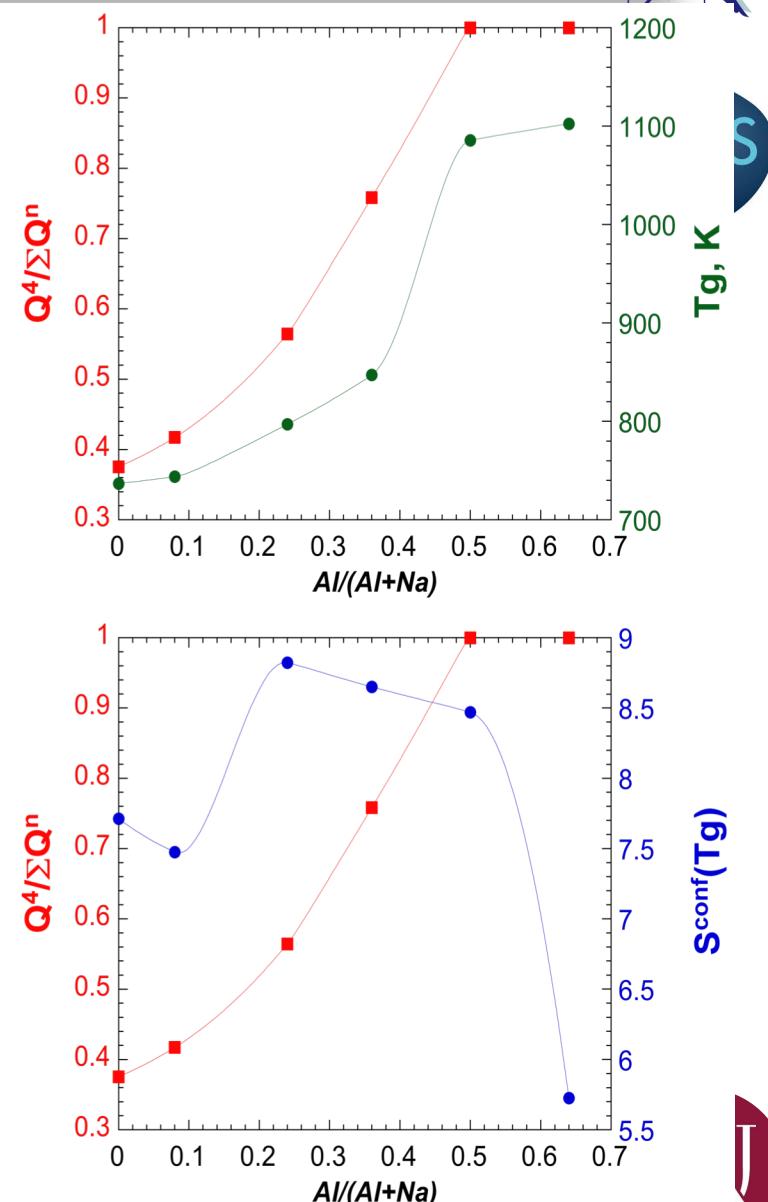
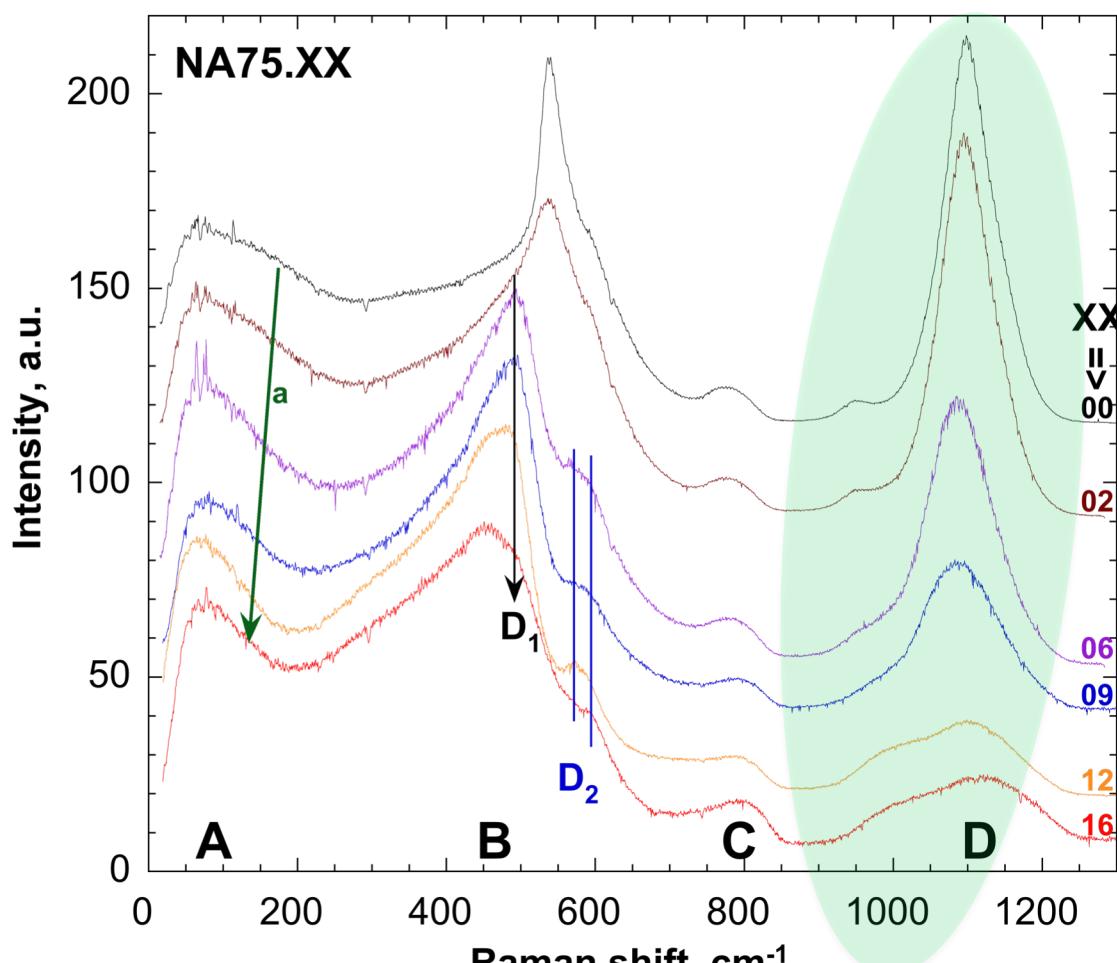
NS3  $\rightarrow$  Albite  $\rightarrow$  NA75.16



$\eta$  increases with  
 $\text{Al}/\text{Na} \leq 1$  and  $\text{Al}/\text{Na} \geq 1$

# $\text{Na}_2\text{O}-\text{Al}_2\text{O}_3-\text{SiO}_2$ Raman spectroscopy

T64000, Jobin-Yvon-Horiba, 488nm

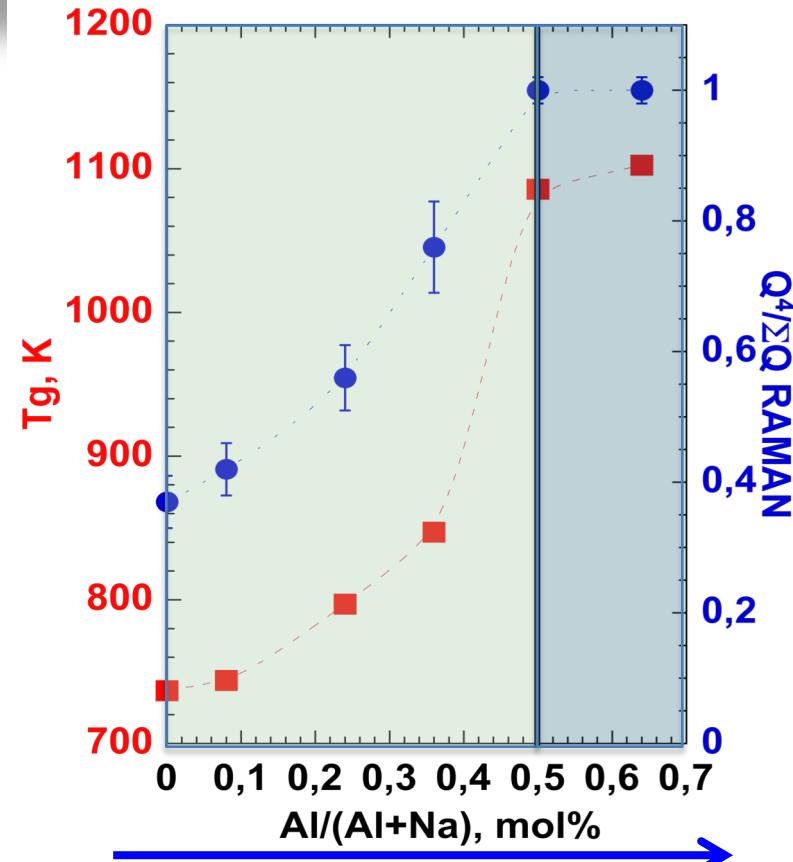
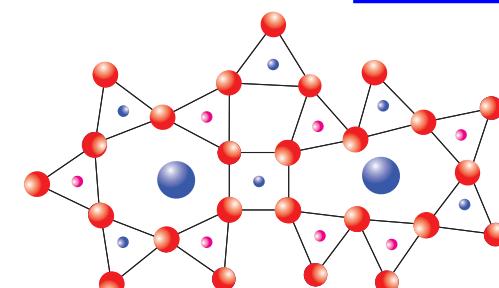
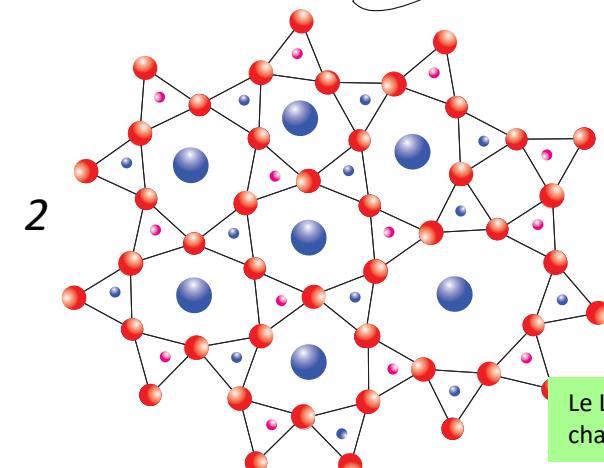
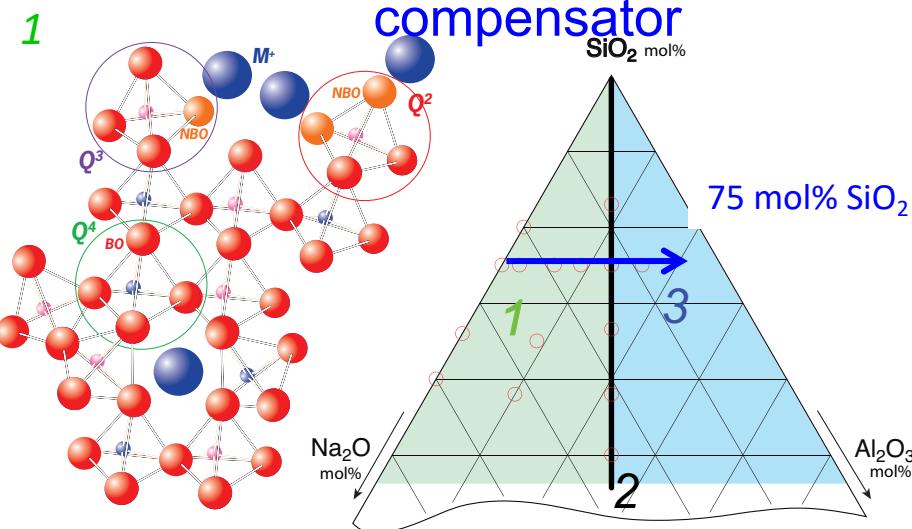


# Summary Al effect

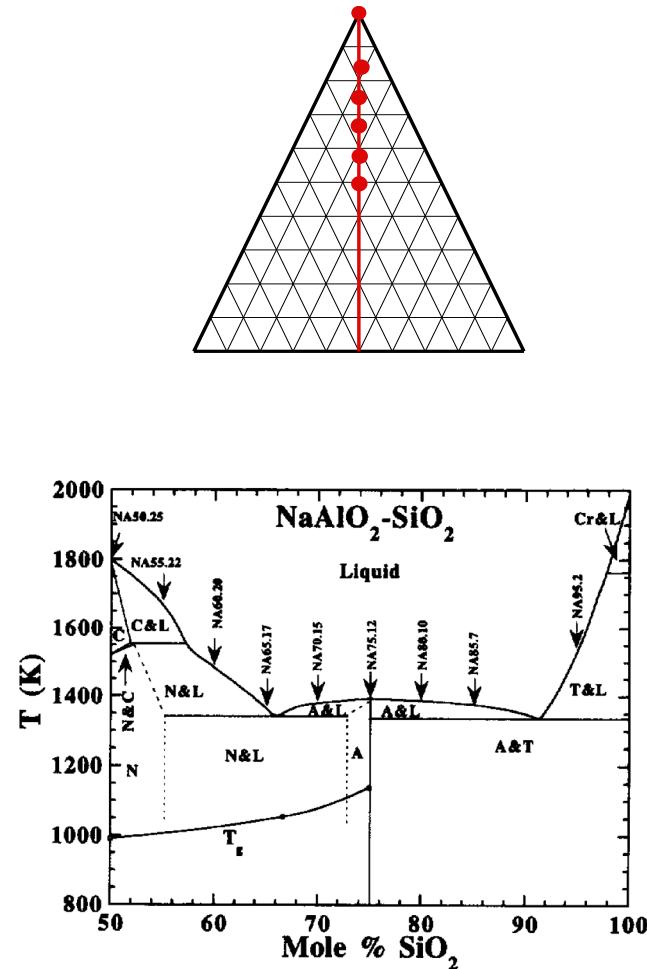
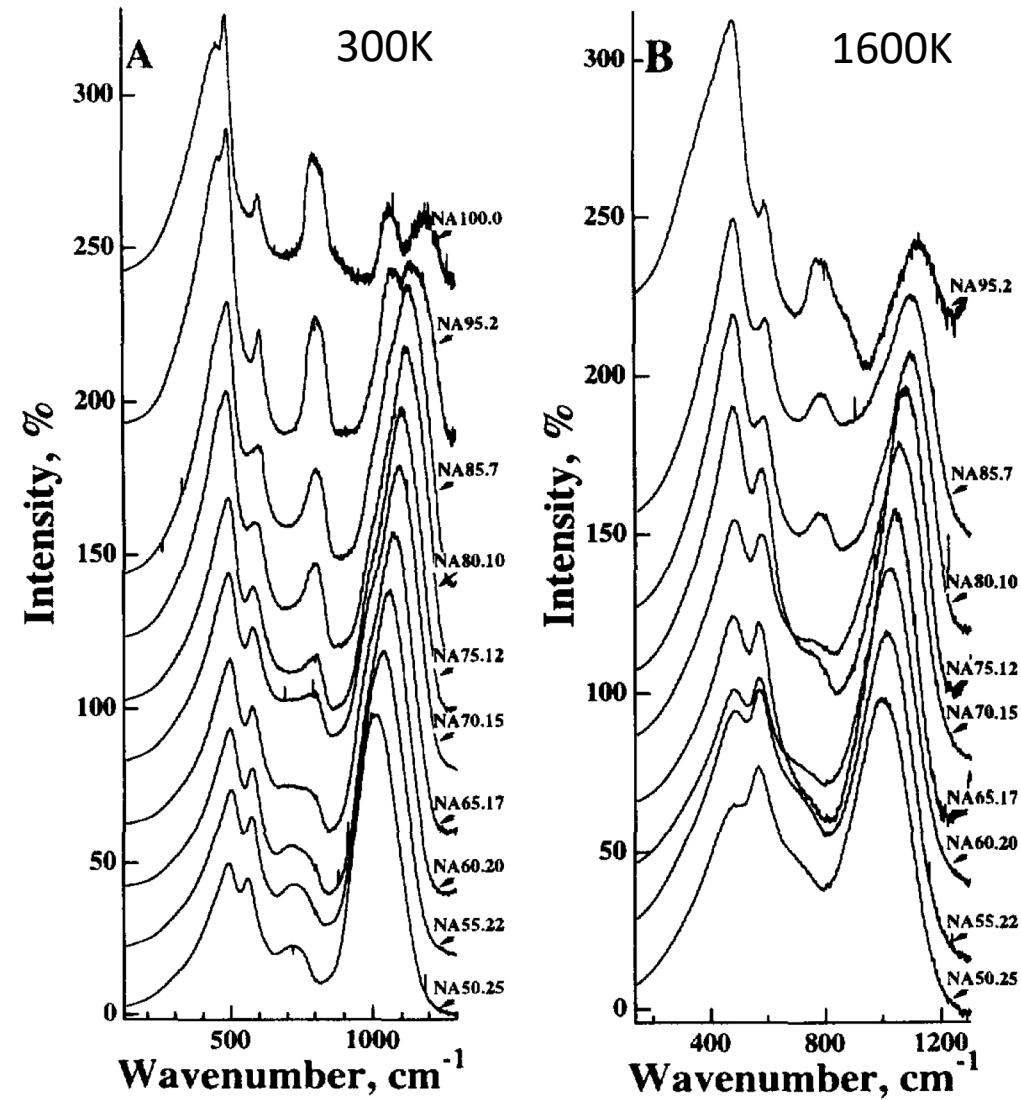
75 mol% SiO<sub>2</sub>

# Geomaterial

- Na<sub>2</sub>O substitution by Al<sub>2</sub>O<sub>3</sub> :
- ⇒ Polymerization
- ⇒ Change Q<sup>3</sup> in Q<sup>4</sup>
- ⇒ Al in Q<sup>4</sup> and Na charge compensator

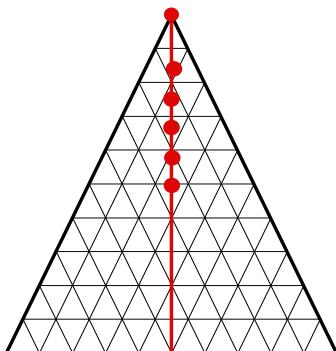
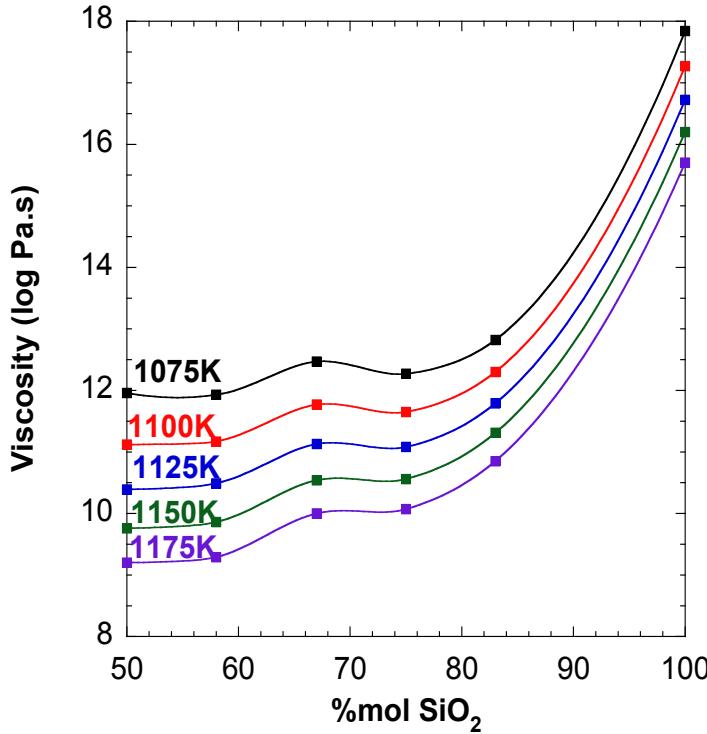
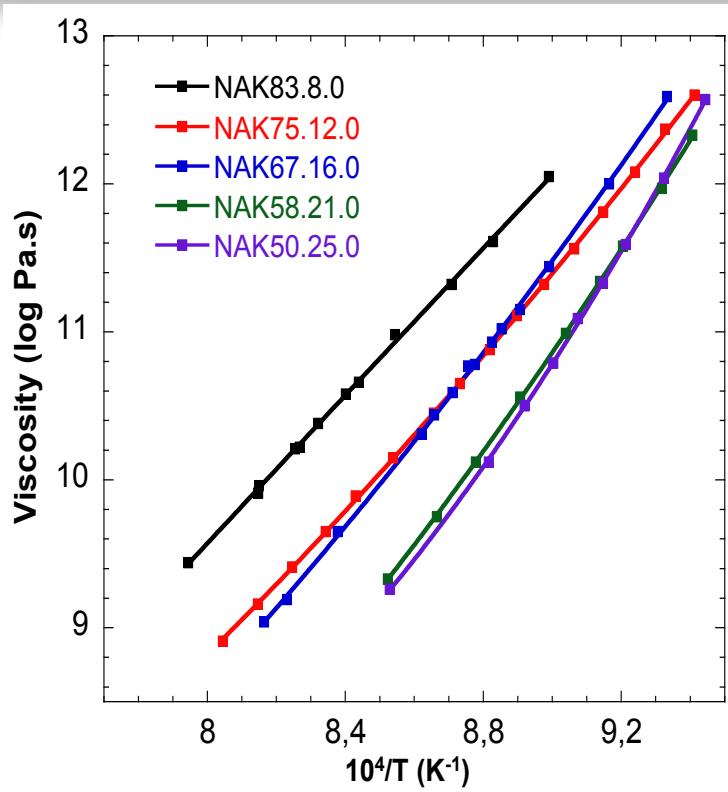


Al in CN 5 in  
peraluminous  
domain (3)



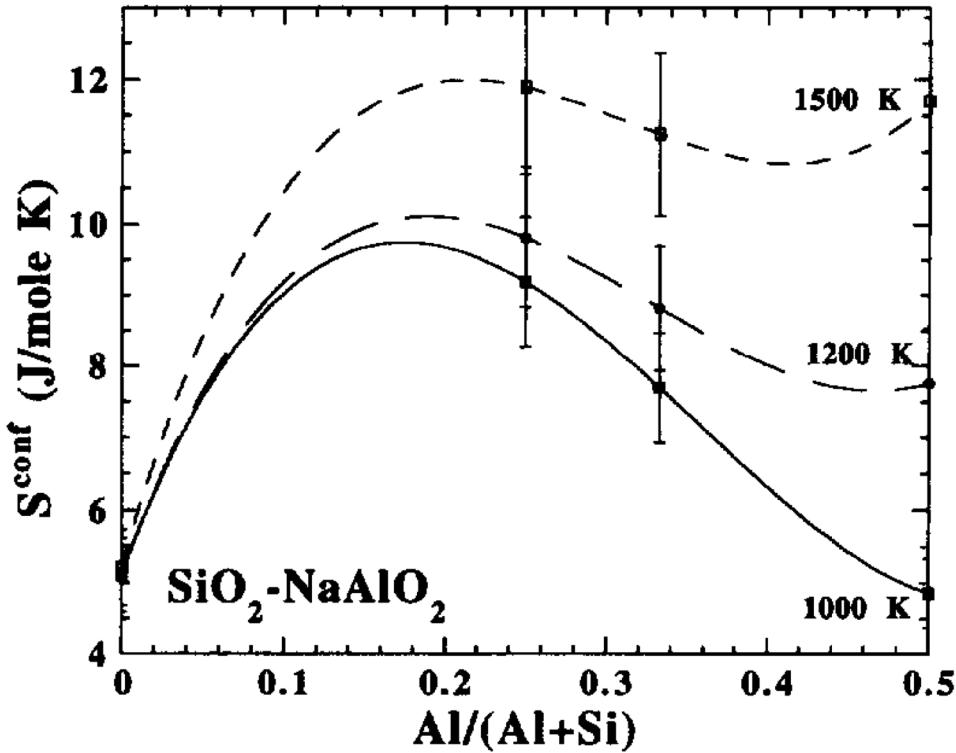
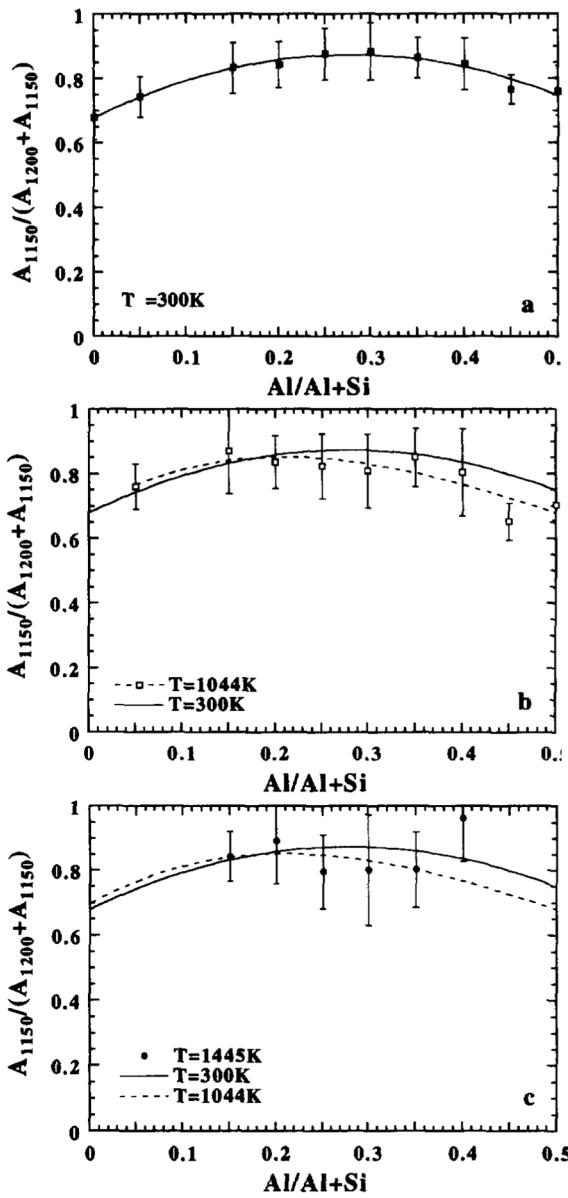
Neuville D.R. and Mysen B.O (1996) Role of aluminium in the silicate network: in situ, high-temperature study of glasses and melts on the join SiO<sub>2</sub>-NaAlO<sub>2</sub>. Geochimica Cosmochimica Acta., 60, 1727-1737





⇒ Variation of viscosity : not a linear function of SiO<sub>2</sub>

⇒ Near the Tg : importance of the configurational entropy



Neuville D.R. and Mysen B.O (1996) Role of aluminium in the silicate network: in situ, high-temperature study of glasses and melts on the join SiO<sub>2</sub>-NaAlO<sub>2</sub>. *Geochimica Cosmochimica Acta.*, 60, 1727-1737



The configurational properties of melts and glasses provide fundamental information needed to characterize magmatic and industrial processes.

The configurational entropy gives a strong idea about glass structure.

It is possible to link the “macroscopic” configurational entropy with the structure of melts determine by NMR or Raman spectroscopy.

Ca/Mg can be mixed randomly in silicate and aluminosilicate glasses and melts

Na/Ca and Na/Sr or Na/M<sup>2+</sup> are not mixed randomly in silicate and aluminosilicate glasses and melts.

Na/K are mixed randomly in silicate and non randomly aluminosilicate glasses and melts. It is possible to determine viscosity from Q species.

Si/Al are mixed randomly in tectosilicate glasses and melts.

Coming soon : a general model between Q species and viscosity

